Jurassic and Cretaceous foreland basin deposits of the Russian Arctic: Separated by birth of the Makarov Basin?

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The age and mode of formation of the various sub-basins of the Amerasian Basin of the Arctic Ocean remain unknown. Jurassic-Cretaceous synorogenic foreland basin deposits are the youngest stratigraphic units deposited in the Russian Arctic prior to rifting and formation of the Amerasian Basin. U-Pb dating of detrital zircon suites (6 samples, ~ 100 zircons each) by LA-ICP-MS reveal that sandstones in the New Siberian Islands have nearly identical sources to those in Chukotka and the South Anyui Zone (SAZ) despite evidence for proximal derivation and little transport. These include abundant Precambrian (~2.1-1.7 Ga), Late Paleozoic (~330-250 Ma) and lesser Mesozoic (~175 to 145 Ma) ages; youngest zircons are likely derived from Main Belt granites in the N. Verkoyansk. The foreland basin and its counterpart the orogenic highlands either extended continuously for ~1400 km along strike, or the localities studied were once much closer together. We hypothesize that rifting/extension associated with formation of the Makarov Basin and development of the SAZ as a transform fault might be one way of explaining the present separation of the study sites

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Introduction

The age of rifting and the mode of formation of the Amerasian Basin of the Arctic have far-reaching implications for the origin of its broad continental shelves, their oil and gas potential, the role the Arctic Ocean has played in global climate, and, most recently, for claims to extend outer continental shelf limits under Article 76 of UNCLOS (e.g., Macnab 2006). Despite the fact that the plate tectonic origin of the Amerasian Basin has been cited as a top research priority by numerous international working groups, logistics and cost continue to preclude scientific drilling of its seafloor. The Gakkel Ridge in the Eurasian Basin is the northern continuation of the mid-Atlantic Ridge (Fig. 1A). Restoration of its Cenozoic spreading history places the Lomonosov Ridge, a strip of continental crust, against the Barents Shelf (e.g., Rowley & Lottes 1988) (Fig. 1B), leaving the older and more controversial Amerasian Basin to be explained by other rift mechanisms (Fig. 1A,B). The Amerasian Basin consists of the Makarov Basin, the Alpha and Mendeleev Ridges, and the Canada Basin (Fig. 1A). The origin of these basins and intervening highs have been debated for years (e.g. review in Lawver & Scotese 1990). Grantz et al. (1979) proposed a rotational opening model for the Amerasian Basin where the Lomonosov margin of the Amerasian Basin operates

as a strike-slip or transform fault boundary, a model which remains the most cited today. In the rotational rift model, the Alpha and Mendeleev Ridges must be parts of a younger feature, perhaps a hot spot related volcanic edifice (e.g., Lawver et al. 2002). Previous interpretations of the Makarov Basin included one which interpreted it as a Late Cretaceous oceanic basin that formed parallel to the Lomonosov Ridge by rift opening in a direction orthogonal to the Lomonosov Ridge based on magnetic anomalies (e.g. Sweeney et al. 1982; Taylor et al. 1981; Vogt et al. 1982). In this model, the Alpha and Mendeleev Ridges with their subparallel basins and highs (Fig. 1A) might represent a component of what was rifted away from the Lomonosov Ridge. More recent magnetic data (Glebovsky et al. 2000) do not reveal the earlier interpreted anomalies, and seismic refraction studies show that the Mendeleev Ridge is underlain by thick (~ 34 Km) but mostly mafic (6.8-7.6 km/sec) crust (e.g., Ivanova et al. 2006). A detailed analysis of the Lomonosov Ridge by Cochran et al. (2006) argues in favor of the Grantz et al. (1979) model. Thus the "rotational rift" model is the most agreed upon model for the formation of the Amerasian Basin. This model makes several specific predictions for the Russian Arctic. The rotational rift model argues that the Arctic Alaska-Chukotka microplate (AACM) rotated from the Canadian Arctic margin to where it is today in Arctic Russia, closing an ocean basin,



Figure 1. A. Index and bathymetry of the Arctic Ocean (IBCAO, 2002) showing location of the Amerasian Basin and its features. Boxes show the approximate locations of Figures 2 (New Siberian Islands) and 3 (Chukotka). B. Schematic diagram showing the proposed size and shape(s) of the Arctic Alaska-Chukotka plate and its inferred restored position along the Canadian Arctic margin prior to rotational rift opening of the Amerasian Basin. Light shaded region (pink) is the location of the inferred Anyui Ocean that once lay between Siberia and North America. Basemap of Figure 1B is from Rowley & Lottes (1988) and shows the Eurasian or North Atlantic basin closed at 140 Ma. Sandstones we compare in this paper are shown at opposite ends of a hypothetical foreland basin, the Myrgovaam (Rauchua) Basin. The South Anyui Zone (SAZ) is shown by a black line that continues into Alaska as the Angayucham zone, and is characterized by remnants of oceanic and island arc terranes. Red stars are loci of Main Belt granitoid plutons studied and dated by Toro et al. (2007), Prokopiev et al. (2007) and Akinin et al. (in press).

the "Anyui Ocean" south of the rotating plate (Fig. 1B) (e.g., Grantz et al. 1979; Grantz & May 1982; Rowley & Lottes 1988; Grantz et al. 1990; Lavwer et al. 2002). The time interval of rotation is 135-120 Ma (Fig. 1B) (Lavwer et al. 2002). Rowley & Lottes (1988) discuss another key geologic aspect of the rotation model: It is viable only if the Arctic Alaska-Chukotka microplate (AACM) is the right shape and size so that it doesn't significantly overlap the Lomonosov Ridge and (once) adjacent Barents Shelf in the process (Fig. 1B). Thus, supporters of the rotation model draw the western end of the rotating plate at some arbitrary position in the Russian Arctic that allows for a better fit (e.g., Rowley & Lottes 1988; Grantz et al. 1979; Grantz & May 1982; Lavwer et al. 2002; Kuzmichev 2009) (Fig. 1B). Problems presented by the rotational opening model for Arctic Russia include the fact that geologists have long cited evidence that the southern boundary of the AACM, the South Anyui suture zone (SAZ) extends to the New Siberian Islands (e.g. Parfenov & Natali'in1985; Fujita & Cook 2000; Kuzmichev et al. 2006; Kuzmichev 2009), making the plate too big to fit the rotation model (Fig. 1). More recently, detrital zircon geochronology of Triassic strata from the margins of the Amerasian Basin provided a very preliminary data set to evaluate this model (Miller et al. 2006). This new data suggested that Chukotka has greater affinity with NE Russia instead of the Canadian margin.

This paper directly addresses the question of the paleogeographic origin of Chukotka, the extent of the AACM and the timing of its deformational events by comparing the tectonic setting, petrography and single-grain U-Pb ages of detrital zircons of Volgian or Tithonian to Valanginian (Late Jurassic to Early Cretaceous) syn-orogenic foreland basin sandstones in two distant parts of the AACM, Stolbovoi Island, (New Siberian Islands) and central Chukotka and the South Anyui Zone (SAZ) (mainland Arctic Russia), localities that are now ~ 1400 apart (Fig. 1). The syn-orogenic sandstones are derived from collision or shorteningrelated orogenic highlands that lay to the south of the SAZ. Foreland basin sediments were shed northward across platformal sequences that underlie the northern part of the New Siberian Island Archipelago and across an actively deforming region of the Chukotka fold belt in Chukotka. Their remarkable similarity in the two regions provides new data with which to constrain plate models for the Amerasian Basin of the Arctic.

Regional geology

Deformation of Paleozoic and Mesozoic sequences of the AACM and closure of the Anyui oceanic basin along its southern edge led to the deposition of the Jurassic-Cretaceous syn-orogenic foreland basin sequences described here. This deformation has been recently studied on Stolbovoi and Big Lyakhov Islands by Kuzmichev et al. (2006) and Kuzmichev (2009), in Chukotka by Katkov et al. (2006), Miller et al. (in press) and Miller & Verzhbitsky (in press), and in the SAZ by Sokolov et al. (2002) and Bondarenko et al. (2003) (Figs. 2 and 3). In these localities, deformation has been linked to crustal shortening and arc collision and lesser strikeslip faulting (Sokolov et al. 2002). As described by these authors, deformation was accompanied by uplift which led to the deposition of southerly-derived, northerlytransported synorogenic clastic sequences. These clastic sequences are mostly basinal gravity flow deposits derived from mixed sources that include deformed platform to basinal Paleozoic-Mesozoic sequences, coarse

Figure 2. Simplified geologic and tectonic map of the New Siberian Islands and the South Anyui Suture Zone (SAZ) after Vol'nov et al. (1998) and Kuzmichev et al. (2005, 2006). Black star is location of sample analyzed for detrital zircons in this study (for exact location see Table 2). Light grey patterns outline the extent of allochthonous arc and oceanic rocks (v pattern), Tithonian to Valanginian syn-orogenic basin deposits (dots) and platformal rocks (limestone pattern). Age of granitic rocks is from Layer et al. (2001), Dorofeev et al. (2001) and unpublished data Rounded rectangles of Kuzmichev. are ⁴⁰Ar/³⁹Ar biotite plateau ages and rectangles are U-Pb zircon ages.



Figure 3. Simplified geologic and tectonic map of Chukotka showing the Anyui-Chukotka fold belt, the distribution of Jurassic-Cretaceous syn-orogenic deposits and the younger (syn-extensional) plutons that intrude the fold belt (with U-Pb ages in squares), after Miller and Verzhbitsky (in press). Stars are locations of detrital zircon samples (for exact localities see Table 2).





Figure 4. Point-count data and photomicrographs of Jurassic-Cretaceous samples. A. Sample ELMCH03.24.3, x-polars, field of view ~ 3mm. Note abundance of angular plagioclase, K-feldspar, monocrystalline quartz and biotite. B. Sample 53.1 from Stolbovoi, x-polars, field of view ~ 1 mm. C. Jurassic-Cretaceous sandstone from Chukotka with abundant rock fragments of deformed slate (most likely from deformed Triassic and older strata), planar light, ~ 2mm across. D. Jurassic-Cretaceous from Chukotka with abundant volcanic detritus including monocrystalline embayed quartz phenocrysts from felsic volcanic sources, planar light, ~ 2mm across.

Table 1. S	Sandston (Point count	e petrogr data from sa	aphy ndstones coll	lected from J	3-K1 deposit	of Russian A	Arctic)
Sample	Q%	F%	L%	Mtx%	Lv	Ls	Lm
50/1-02 Stolbovoi	30.3	56.3	13.3	26.6	55.6	33.3	11.1
51/1-02 Stolbovoi	34.3	51.3	14.3	24.4	52.0	32.0	16.0
53/1-02 Stolbovoi	31.3	49.8	18.9	28.5	31.6	63.2	5.3
54/1-02 Stolbovoi	30.7	52.0	17.3	23.3	48.3	37.9	13.8
CH03ELM 24.3	9.35	61.94	28.75	0.92	82.02	12.36	5.6
CH04 JT54A	31.75	33.33	34.92	9	63.64	18.2	18.2
CH04 JT54D	14.8	52.8	32.4	4.2	67.9	19.75	12.35
CH04 ELM7	18.8	41.03	40.17	1.63	72.34	15.96	11.7
CH04 ELM23A	14.11	28.22	57.67	15	73.4	25.53	1.06
GB9986 SAZ	33.47	26.45	40.08	3.2	29.9	38.14	15.5
GB9947 SAZ	23.81	26.98	49.21	4.25	83.9	12.9	3.23
9947/1 SAZ	32.73	38.2	29.1	13.3	87.5	12.5	0
Notes: Q%=0 L _s =Ls/(Lv+Ls	Q/Σ; F%=F/Σ s+Lm), L _m =L	C; L%=(Lv+Ls m/(Lv+Ls+L	s+Lm+Col)/2 m) v - volcar	E; Mtx%=Mt nic, s - sedim	x/(Mtx+Σ); L entary, m – n	v=Lv/(Lv+Ls netamorphic	+Lm),

crystalline rocks, and volcanic/plutonic rocks. In the New Siberian Islands, Jurassic-Cretaceous syn-orogenic strata are interpreted to have been deposited northwards from a region of allochthonous arc and ophiolitic rocks exposed on Big Lakhov Island and the Russian mainland (Kuzmichev et al. 2006) (Fig. 2). In Chukotka, the map distribution of Jurassic-Cretaceous strata extends at least as far south as the SAZ (Fig. 3), where deformed strata of this age underlie most of that zone. Despite the complex deformation in this zone, Jurassic-Cretaceous syn-orogenic sandstones are inferred to have been derived from the south (e.g., Sokolov et al. 2002; Bondarenko et al. 2003). The syn-orogenic sediments of Chukotka were deposited across the evolving Chukotka fold belt, and are in turn deformed as well. The northernmost exposures are found along the Arctic coast near the town of Pevek (Fig. 3), where they have much lower sand:shale ratios than strata further south. Sparse fossils from all of these localities indicate Late Jurassic to Early Cretaceous ages (Volgian or Tithonian to Valanginian (e.g. Samusin & Belousova 1985; Paraketsov & Paraketsova 1989), which, for the sake of convenience, we refer to as "Jurassic-Cretaceous" (J3-K1 in Russian literature). Within the SAZ proper, Dovgal (1964) describes strata with fossils as young as Hauterivian-Barremian that are involved in at least the latest stages of deformation (see also summary in Bondarenko et al. 2003). The deformed Jurassic-Cretaceous syn-orogenic deposits are intruded by post-tectonic plutons at least as old as 121 Ma in the New Siberian Islands (Fig. 2) (Layer et al. 2001; Dorofeev et al. 2001; Kuzmichev 2009) and at least as old as 116.9 ± 2.5 Ma in Chukotka (Fig. 3)(Katkov et al. 2007; Miller et al. in press). Quaternary deposits cover most of the topographically low coastal region of the Russian Arctic that lies between Chukotka and the New Siberian Islands (Fig. 1A), so there is no map continuity between any of the units exposed in these two distant locations.

Sandstone description and petrography

New Siberian Islands

The New Siberian Islands are represented mostly by a Paleozoic-Mesozoic platformal sequence (Kos'ko et al. 1985, 1990; Fujita & Cook 1990; Dorofeev et al. 1999) which, prior to opening of the Eurasian basin, was attached to the Siberian edge of the Barents Shelf (e.g. Rowley & Lottes 1988; Kuzmichev & Pease 2007) (Fig. 2). The three southern islands of the New Siberian Island archipelago are underlain mostly by Jurassic-Cretaceous syn-orogenic sandstones and shales that were shed across the southern margin of the New Siberian platform when it was overridden by allochthonous arc and oceanic rocks of the Anyui Zone (Fig. 2) (Kuzmichev et al. 2005, 2006; Kuzmichev 2009). The syn-orogenic sandstones thus constitute the deposits of a foreland basin that lay to the north and east of the zone of thrust-related deformation. The syn-orogenic foreland basin deposits can be classified into a lower and an upper part representing the initial and the main stages of basin formation.

The first stage of deposition is represented by synorogenic deposits in southeastern Big Lyakhov Island where they are partly tectonically buried under the allochthonous ophiolitic assemblage (Kuzmichev et al. 2005; Kuzmichev 2009). They differ from the laterstage deposits by their sandstone composition and more extensive involvement in deformation. The age of these deposits is not reliably confirmed by fossils. They are thought to be Late Jurassic because they contain ophiolitic debris and the ophiolites have been dated as Jurassic (most likely Oxfordian-Kimmeridgian) (Kuzmichev & Lebedev 2008). Triassic-age sandstones and shales are also known to be present in this more highly deformed region (Kuzmichev et al. 2006).

Clastic successions representing the main stage of basin formation crop out in western and northern Big Lyakhov Island and underlie most of Small Lyakhov Island and the Stolbovoi Island (Fig. 2). In the first location they are contact metamorphosed by nearby plutons and were referred to as Volgian only by analogy with similar fossiliferous deposits on Stolbovoi and Small Lyakhov islands (Samusin & Belousova, 1985). On Small Lyakhov Island, Mesozoic sediments are locally exposed at the northern and eastern edges of the Island and were penetrated in several drill-holes in the central part of the Island. On the northeastern shore of the Island, they contain *Buchia* of Volgian-Valanginian age. The same deposits are best exposed on Stolbovoi Island, which is almost completely surrounded by sea cliff exposures.

Stolbovoi Island was visited twice by one of us (Kuzmichev). In 2002, sea cliff exposures on the northwest coast of the island were studied, and the sample discussed in this paper, collected. In 2007, the southern half of the island was investigated in detail and these results are summarized below.

A sand-rich distal turbidite complex underlies Stolbovoi Island. The rocks are slightly altered but preserve all their original sedimentary features. The strata are gently folded and cut by faults. A notable feature of the turbidite complex is the presence of light-colored, mediumgrained, amalgamated sandstone units up to 25 meters thick. The detrital zircon sample, discussed below, was obtained from one of these thick sandstone beds. The rest of the complex consists mostly of mud-rich rhythmically bedded, dark-colored sandstones which contain rip-up clasts and plant debris. Shale-silt dominated intervals with minor sandstones with abundant ripples represent a lesser part of the sequence. The massive sand units are interpreted as redeposited in deep water after accumulation on an outer shelf environment. The dark sandstones and shales are interpreted as sediments directly transported from land across a narrow shelf into deep water in the manner described by Mutti et al. (2003). The third facies represents the final deposition of sediments from suspension after the main turbidity current waned.

Sole marks such as tool marks, groove casts and flute casts indicate that sediments were transported from south to north (the mean of 78 measurements is 357°). Ripple marks indicate current directions from the southwest towards the northeast (the mean of 427 measurements is 56.5°). *Buchia* fossils were found in more than 20 locations in the studied area. They were identified by V. A. Zakharov and in general indicate a Late Volgian (Tithonian) to Early Valanginian age.

Chukotka

Jurassic-Cretaceous strata in exposures north of Bilibino to Pevek on the coast (Fig. 3) were referred to as the Raucha sedimentary basin by Nockleberg (1994, 1998) after Paraketsov & Paraketsov (1989), and as the Myrgovaam Basin by Baranov (1996). Sandstones from the South Anyui Zone have been described in detail by Bondarenko et al. (2003) so our description below focuses on the nature of exposures north of Bilibino.

Most of these sediments are massive, well-cemented, resistant arkosic sandstones that represent gravity flow deposits (probably grain flow and/or amalgamated Bouma A horizons) and contain Volgian (Tithonian) to Valanginian fossils at scattered localities (Akimenko 2000; Belik & Sosunov 1969; Sosunov & Tiliman 1960). Sequences exposed along the coast near Pevek are thinner-bedded gravity flow deposits with lower sand: shale ratios.

The Jurassic-Cretaceous sediments of the Myrgovaam Basin are often described as unconformable on older rocks on map legends and cross-sections. Nockleberg et al. (1994, 1998) classified the sediments as "post-tectonic". Baranov (1996), however, describes them as structurally detached and imbricated by N-verging thrust sheets. Our field studies (2003, 2004) support Baranov's conclusions: underlying, Upper Triassic to Lower Jurassic thin bedded shales and sandstones are tightly folded while overlying, massively bedded Jurassic-Cretaceous sandstones form large, coherent fault-imbricated panels. Conglomerates are not found along basal contacts of the overlying massively bedded deposits, but fault zone gouge and breccia are common, suggesting slip along a fault contact. Overlying beds often dip gently into and are truncated by the basal contact, also supporting structural discordance rather than an unconformity (Baranov 1996; Miller et al. 2006). Within

strata mapped as Jurassic-Cretaceous near the Rauchua River, we observed thin-bedded arkosic sandstones in conformable stratigraphic sequence with underlying Jurassic shales. Map-scale conglomeratic lenses, inferred to be submarine channel deposits, cut into these thinbedded strata and all strata are deformed together in N-vergent open to tight folds (Miller et al. 2006). These relations suggest original stratigraphic continuity between underlying thinner-bedded sequences and overlying massive sandstones. Near Pevek, Jurassic-Cretaceous strata are involved in folds with a south-dipping axial planar cleavage (Miller & Verzhbitsky in press), and thus are clearly involved in at least the latest part of the shortening history of Chukotka.

Poor sorting and angular grains together with fresh plagioclase, K-feldspar, biotite and muscovite point to proximal orogenic sources that contained granitoid/ gneissic rocks in addition to sedimentary, meta-sedimentary and volcanic rocks that are present as abundant lithic fragments in these rocks (Table 1) (Fig. 4). Volcanic lithic fragments are dominated by variable amounts of intermediate to felsic volcanic rocks and rarer mafic volcanic rocks (Fig.4) (Table 1). Sedimentary and low-grade metasedimentary rocks also form an important component of the sandstones (Table 1). These are mostly fragments of shale, siltstone and multiply deformed slate and phyllite (Fig. 4). Heavy minerals include variable percentages of zircon, apatite, and pyrite, magnetite and ilmenite.

Detrital zircon data

Detrital zircons from six samples were dated by LA-ICPMS (approximately 100 single grain ages per sample). Sample locations are listed in Table 2, description of analytical methods in Appendix 1, and measured isotopic ratios and interpreted ages in Table 3. The U-Pb ages from each sample are plotted on relative age probability distribution diagrams [Ludwig, 2003] (inset, Fig. 5) and in Figure 6. The Kolmogorov-Smirnov (KS) statistical test was used to further assess the similarity of the distributions of single grain ages and these results are presented in Table 4.

Cumulative age-probability plots (Gehrels 2006), provide the best comparison between data from individual

Table 2. Location	ons of samples o	analyzed for det	rital zircon		
Field Number	Area	Lat	Long	Age	Map Unit
GB9986	South Annui Zone	67°17'20"	164°47'13"	J3-K1	(Bondarenko et al., 2003)
CH04 ELM7	Chukotka	68°29'29"	165°41'27.2"	J3-K1	Pogynden Fm.
04JT54C	Chukotka	68°25'19"	166°40'14"	J3-K1	Pogynden Fm.
ELM CH03.24.3	Chukotka	69°09'45.6"	165°02'57.7"	J3-K1	Pogynden Fm.
ELM06 PV10	Pevek, Chukotka	69°43'34.4"	170°57'8.1"	J3-K1	Rauchua Basin, Chauna Fm.
53-1-02	Stolbovoi Is.	74°11'45.3"	135°27'44.9"	J3-K1	Stolbovskaya Fm.

Table III. U-I	^o b (zirco	n) geod	Irono	logic and	alyses	by Laser	Abla	tion Mu	Iticollect	for ICP	Mass S _I	oectror	netery				
					lsotopic	ratios					Apparent a	ges (Ma)					
Analysis	U (ppm)	206Pb	U/Th	207Pb*	+I	206Pb*	+	error	206Pb*	+	207Pb*	+1	206Pb*	+1	Best age	+	Age of youngest cluster (Ma)
		204Pb		235U	(%)	238U	(%)	corr.	238U	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(3+ analyses)
																	145 (3)
Sample GB9986 5	outh Annyu	ii Zone															164 (3)
GB9986-47	136	2214	1,7	0,1467	8	0,02282	1,9	0,24	145,5	2,7	139,0	10,3	29,5	185,3	145,5	2,7	
GB9986-42	182	862	1,1	0,18368	18,8	0,0234	2,8	0,15	149,1	4,1	171,2	29,7	489,1	414,1	149,1	4,1	
GB9986-8	152	541	0,6	0,23799	40,9	0,02453	5,3	0,13	156,2	8,2	216,8	79,9	939,1	865,7	156,2	8,2	
GB9986-24	116	1642	1,6	0,16528	15,8	0,02562	1,8	0,11	163,1	2,9	155,3	22,8	38,7	377,5	163,1	2,9	
GB9986-103	27	614	1	0,1684	12,7	0,02588	2,5	0,2	164,7	4,1	158,0	18,6	59,4	297,7	164,7	4,1	
GB9986-73	107	1639	0,9	0,16719	15,8	0,0261	2,3	0,14	166,1	3,7	157,0	23,0	21,2	377,1	166,1	3,7	
GB9986-6	586	11572	1,9	0,18473	1,7	0,02726	-	0,59	173,3	1,7	172,1	2,7	155,3	32,0	173,3	1,7	
GB9986-106	144	3820	0,9	0,1764	7,1	0,02758	1,5	0,21	175,4	2,5	165,0	10,8	17,9	166,3	175,4	2,5	
GB9986-75	212	4411	1,9	0,18236	4,8	0,02766	c	0,63	175,9	5,2	170,1	7,5	90,2	88,9	175,9	5,2	
GB9986-53	356	1371	1,6	0,27234	19,7	0,02945	2,8	0,14	187,1	5,2	244,6	42,8	839,9	409,4	187,1	5,2	
GB9986-64	00	1524	<u>ν, </u>	0,22/04	20,2	0,03533	4,4	0,22	223,8	9,7	207,8	37,9	29,0	4/0,1	223,8	9,7	
GB9986-56	139	4057	0,9	0,23281	12,1	0,03662	3,4	0,28	231,9	7,8	212,5	23,2	3,1	280,5	231,9	7,8	
GB9986-25	39	1110	- '	0,23828	18,5	0,03748	3,1	0,17	237,2	7,3	217,0	36,2	3,0	442,6	237,2	7,3	
GB9986-88	568	15023	1,4	0,27206	1,4	0,03842	-	0,7	243,0	2,4	244,3	3,1	257,2	23,6	243,0	2,4	
GB9986-102	48	1739	0,6	0,25749	17,4	0,0401	1,8	0,11	253,5	4,5	232,6	36,1	27,0	416,5	253,5	4,5	
GB9986-92	71	1882	2,1	0,25582	13,7	0,04018	2,3	0,17	253,9	5,7	231,3	28,4	7,1	326,9	253,9	5,7	
GB9986-105	152	5467	1,2	0,27841	4,1	0,0408	-	0,25	257,8	2,5	249,4	9,0	171,0	92,1	257,8	2,5	
GB9986-62	57	1761	0,5	0,27076	15,5	0,0422	1,4	60'0	266,5	3,7	243,3	33,4	25,2	371,0	266,5	3,7	
GB9986-23	1228	33907	20,6	0,30144	1,4	0,04261		0,71	269,0	2,6	267,5	3,3	254,8	23,1	269,0	2,6	
GB9986-18	635	9582	2,4	0,3162/	2,5	0,04294	6,1	0,74	2/1/0	5,0	279,0	6,2	346,5	38,4	2/1/0	5,0	
GB9986-81	178	4854	1,3	0,30066	3,9	0,04368	-	0,26	275,6	2,7	266,9	9,2	191,2	88,1	275,6	2,7	
GB9986-44	103	3297	0,9	0,31956	9	0,04686	1,9	0,32	295,2	5,5	281,6	14,8	169,5	133,4	295,2	5,5	
GB9986-95	116	3447	0,9	0,3199	5,8	0,04687	1,1	0,19	295,3	3,2	281,8	14,2	171,4	132,4	295,3	3,2	
GB9986-99	190	6659	1,4	0,3329	3,6	0,04762	1,1	0,31	299,9	3,3	291,8	9,2	227,4	79,3	299,9	3,3	
GB9986-108	242	9580	3,3	0,34413	1,9	0,04777		0,52	300,8	2,9	300,3	11.7	296,4	38,0	300,8	2,9	
00-004605	061	01 00	0,0	0,204.00	t 1	0,00400	- 0	07'0	7170	0,0	41000	7/11	1,002	0//0	041,Z	0,0	
GB9986-59	148	8936	12.8	0.61358	0,0	0.07988	<u>יא</u>	0,29	495.4	4.8	241,0 485.8	19,1	C,222 440.8	55.9	495.4	0,1	
GB9986-38	59	3869	0.7	0,88098	6,4	0,10918	2.7	0.42	668,0	16,9	641,5	30,4	549,3	126,7	668.0	16,9	
GB9986-57	102	10912	1,1	1,35739	1,8	0,14374	1,4	0,82	865,8	11,7	870,7	10,3	883,4	21,1	865,8	11,7	
GB9986-110	212	38602	2,8	1,73346	1,5	0,17205	-	0,68	1023,4	9,5	1021,0	9,5	1016,0	21,9	1016,0	21,9	
GB9986-58	152	33922	1,8	5,24523	1,4	0,33377	-	0,71	1856,6	16,1	1860,0	12,1	1863,8	18,1	1863,8	18,1	
GB9986-17	376	51988	2,9	4,07792	1,5	0,25914	1,1	0,73	1485,4	14,3	1649,9	12,0	1866,2	18,1	1866,2	18,1	
GB9986-10	60	19309	2	4,72569	1,6	0,3002	1,2	0,78	1692,3	18,5	1771,8	13,4	1866,8	18,1	1866,8	18,1	
GB9986-22	179	38171	5,4	4,97025	1,8	0,31561	1,1	0,62	1768,2	17,4	1814,3	15,5	1867,5	26,0	1867,5	26,0	
GB9986-109	337	102143	2,3	5,25163	1,7	0,33267	-	0,58	1851,3	16,1	1861,0	14,7	1871,9	25,3	1871,9	25,3	
GB9986-14	47	9779	1,4	5,57433	2	0,35178	4,9	0,98	1943,1	82,4	1912,1	43,2	1878,7	18,2	1878,7	18,2	
GB9986-79	397	68076	2,3	4,48327	1,4	0,28098	-	0,71	1596,3	14,1	1727,9	11,7	1891,2	18,0	1891,2	18,0	
GB9986-61	216	46985	2	5,30987	1,4	0,33274	-	0,71	1851,7	16,1	1870,5	12,1	1891,4	18,0	1891,4	18,0	
GB9986-83	314	27072	4,1	4,69086	4,9	0,29393	4,6	0,94	1661,2	68,1	1765,6	41,4	1891,5	30,1	1891,5	30,1	
GB9986-26	186	37074	2,2	5,35567	1,4	0,33514		0,71	1863,2	16,2	1877,8	12,1	1893,9	18,0	1893,9	18,0	
GB9986-3	38	10529	0,9	777277	1,0 1	0,34389		0,64	1905,3	C,01	1903,0	13,4	2,0091	C,12	2 5001	C,12	
GB9986-84	33	5322	1,2	5,16903	3,1	0,32101	2,9	0,94	1794,7	45,6	1847,5	26,2	0/2021	18,1	0,cu~1 9,7001	18,1	

	Age of youngest cluster (Ma)	(3+ analyses)																																												
	+1	(Ma)	20,6	21,0	17,9	17,9	26.9	22.8	22,8	19,9	20,3	17,9	17,9	18,8	714	17.9	25,4	22,2	46,0	19,7	24,6	18,8	26,3	17,9	17,9	19,5	17,8	28,5	17,8	30.1	24.5	22,2	18,5	24,3	30,5	42,1	35,8	17,7	75 0	73.8	777	25.0	17,8	23,2	43,3	23,4
-	Best age	(Ma)	1909,6	1913,2	1913,6	1915,4	1910/1	1921.9	1922,8	1923,9	1925,8	1925,9	1926,5	1926,7	1933 9	1936.8	1937,0	1938,0	1938,1	1940,3	1940,8	1946,1	1947,3	1949,3	1954,6	1955,3	1960,0	1971,7	1975,2	1976,2	1987.0	1989,8	1990,9	1996,7	2008,3	2012,0	2015,2	2023,8	2026,4	70802	2469.0	2485.0	2518,0	2525,9	2566,9	2584,2
	+1	(Ma)	20,6	21,0	17,9	17,9	090	22.8	22,8	19,9	20,3	17,9	17,9	18,8	714	17.9	25,4	22,2	46,0	19,7	24,6	18,8	26,3	17,9	17,9	19,5	17,8	28,5	17,8	301	24.5	22,2	18,5	24,3	30,5	42,1	35,8	17,7	1/,1	0'02	2/27	25.0	17,8	23,2	43,3	23,4
-	206Pb*	207Pb*	1909,6	1913,2	1913,6	1915,4	1910/1	1921.9	1922,8	1923,9	1925,8	1925,9	1926,5	1926,7	1933.0	1936,8	1937,0	1938,0	1938,1	1940,3	1940,8	1946,1	1947,3	1949,3	1954,6	1955,3	1960,0	1971,7	1975,2	1976,2	1987.0	1989,8	1990,9	1996,7	2008,3	2012,0	2015,2	2023,8	2026,4	70800	2469.0	2485.0	2518,0	2525,9	2566,9	2584,2
ges (Ma)	+1	(Ma)	13,1	13,5	24,2	12,2	15,6	14.0	30,0	12,9	13,1	12,2	12,2	13,9	13.5	12.2	15,2	17,6	24,2	12,9	23,9	12,6	20,8	12,3	12,3	13,8	12,2	16,9	12,3	12,8	15.0	13,9	13,9	21,0	17,5	47,6	79,0	12,4	12,4	16.1	17.6	18.9	13,6	18,7	56,1	18,8
Apparent a	207Pb*	235U	1918,8	1786,5	1854,0	1923,4	1950 5	1935.7	1937,4	1922,0	1938,3	1929,2	1937,0	1015 0	1938.5	1918,8	1956,4	1883,5	1923,4	1974,0	1956,2	1948,7	1943,0	1958,0	1946,9	1940,2	1945,6	1987,9	1964,0	1951,5 1967 6	2012.4	1967,3	1992,5	2015,0	2004,9	1974,7	1698,0	2003,3	1999,5	2016 7	7343.1	2505.2	2521,3	2459,4	2447,6	2552,1 2610,3
	+I	(Ma)	16,7	16,1	41,7	16,7	170	16,8	54,2	16,7	16,8	16,7	16,8	20,6	16.8	16.5	17,5	26,2	18,8	17,2	40,5	16,8	31,9	16,9	16,7	19,3	16,7	19,0	16,9	706	17.8	16,8	20,8	34,3	17,8	82,5	121,6	17,1	1 / ,0	20.0	18.7	29.1	20,9	29,4	106,3	30,1
	206Pb*	238U	1927,2	1680,0	1801,4	1930,9	1978 1	1948.6	1951,0	1920,2	1950,1	1932,1	1946,8	1996,2	1942 9	1902.1	1974,8	1834,4	1909,7	2006,2	1970,8	1951,2	1938,9	1966,2	1939,5	1926,0	1932,1	2003,5	1953,4	1928,2	2037.2	1946,0	1994,1	2032,9	2001,5	1939,2	1453,1	1983,5	70216	1946 2	22013	2530.2	2525,5	2379,7	2306,5	2511,8 2629,1
	error	corr.	0,66	0,68	0,94	0,71	0.55	0.62	0,93	0,67	0,66	0,71	0,71	0,75	0.64	0.71	0,59	0,8	0,4	0,67	0,87	0,69	0,79	0,71	0,71	0,73	0,71	0,57	0,71	0,68	0.6	0,63	0,76	0,82	0,52	0,9	0,98	0,71	0,/1	0,00	0.57	0.68	0,69	0,73	6'0	0,72
	+1	(%)	1	1,1	2,7		- -		3,2	-	1	1	-	1,2		-	-	1,6	1,1	-	2,4	-	1,9	-	-	1,2	-	1,1	-		ų (-	-	1,2	2	-	4,9	9,3	-		- (ų –	1.4		1,5	5,5	1,4
ratios	206Pb*	238U	0,34846	0,29773	0,32239	0,34923	0,35914	0.35294	0,35345	0,34699	0,35325	0,34949	0,35255	0,36296	0.35173	0,34322	0,35844	0,32919	0,34479	0,36508	0,35761	0,35347	0,35091	0,35663	0,35104	0,3482	0,34948	0,3645	0,35394	0,34867	0.37166	0,35239	0,36253	0,37074	0,36409	0,35097	0,25284	0,36027	0,35818	0 35243	0.40702	0.48066	0,47958	0,4465	0,43018	0,47646 0,50358
lsotopic	+1	(%)	1,5	1,6	2,8	1,4	- v (0	1.6	3,5	1,5	1,5	1,4	1,4	1,6	1 9	1.4	1,8	2,1	2,8	1,5	2,8	1,5	2,4	1,4	1,4	1,6	1,4	1,9	1,4	1,5 0 1	1.7	1,6	1,6	2,4	2	5,5	9'6	1,4	1,4	- 2 2 2	1.9	2	1,5	2	6,1	1,9
	207Pb*	235U	5,61742	4,80903	5,20857	5,64787 5 54102	5 87758	5.72859	5,73997	5,63824	5,74615	5,68546	5,7373	5,90727	5 74755	5.6174	5,86726	5,39123	5,64736	5,98706	5,86619	5,81539	5,77696	5,87798	5,80296	5,75824	5,79483	6,08352	5,91898	5,83416	6.25632	5,94157	6,11608	6,27499	6,2029	5,99187	4,32432	6,19176	6,1651 6 44545	6 28763	9.05007	10.78956	10,97883	10,26958	10,1396	11,34675
	U/Th		1,6	30,2	2	ς α	7 1	1.5	2,2	2,4	3,5	0,7	50,4	2,9	94	4.2	2,8	1,3	29,8	9,9	8,5	9	5,3	5,7	3,5	2,3	6,4	2,9	6,4	1,3	0.7	e co	1,8	5	1,4	2,4	2,6	1,4	3,3	2,4	6,1	1.9	2,6	2,3	5,1	1,3
	206Pb	204Pb	71378	119208	73771	52468 28446	110343	21704	42095	18418	91758	34289	207975	22854	20272	40128	119224	40463	155999	68792	10094	129251	39603	129605	74160	94094	50803	53259	199948	59899 99478	17865	14810	103016	78599	93207	5196	13509	27305	49345 27870	7816	192851	48029	105849	56363	23121	25538 162075
	U (ppm)		315	666	370	231	570	81	165	78	420	105	1274	154	00	195	630	211	757	374	35	601	187	650	582	492	390	169	1081	320	91	263	757	414	365	82	156	101	210	478	782	144	558	204	172	570
	Analysis		GB9986-29	GB9986-72	GB9986-63	GB9986-65	GR0086-08	GB9986-91	GB9986-7	GB9986-20	GB9986-77	GB9986-107	GB9986-33	GB9986-13	GR9986-37	GB9986-82	GB9986-96	GB9986-94	GB9986-78	GB9986-60	GB9986-71	GB9986-28	GB9986-74	GB9986-50	GB9986-49	GB9986-70	GB9986-97	GB9986-101	GB9986-2	GB9986-43 GR0086-34	GB9986-100	GB9986-54	GB9986-16	GB9986-9	GB9986-80	GB9986-36	GB9986-55	GB9986-1	GB9986-85 GB0086-30	GR0086-5	GR9986-32	GB9986-48	GB9986-93	GB9986-69	GB9986-21	GB9986-31 GB9986-4

	Age of youngest cluster (Ma)	(3+ analyses)											157 (6)																																	
	+1	(Ma)	20,5	52,2	1/,8	23.7	16,6	19,0	16,5	16,5	16,4	15,9		3,0	6,3	14,7	8,5	6,9	5,1	/,3	4,8	0,0	69	43	46	0, 8, F 8	5,5	4,1	3,0	15,4	5,4	د 0 ا	40	11,9	9,7	9,0	38,0	22,6	22,1	19,6	34,7	23,4	32,8	18,0	1,22	21, ^y 18,1
-	Best age	(Ma)	2615,0	2629,7	2643,5	2668.4	2673,3	2683,3	2695,8	2737,0	2746,1	3096,3		156,9	158,7	160,8	161,6	163,4	170,5	6'1/1	233,3	241,0	246.1	2481	240,5 248,8	253.4	256,8	256,9	262,2	267,0	///07	0,020	350.6	365,3	370,4	445,0	1735,7	1736,2	1749,1	1834,7	1858,1	1884,1	1888,4	1894,9	1006 0	1906,9
	+1	(Ma)	20,5	52,2	1/,8	23.7	16,6	19,0	16,5	16,5	16,4	15,9		181,3	1006,2	1354,1	1431,4	1243,3	374,6	497,8	2//3	7 6 2 1	3114	7741	2180	659.7	101,5	43,1	60,7	131,4	132,1	2,000	7475	223,2	208,2	140,9	38,0	22,6	22,1	19,6	34,7	23,4	32,8	18,0	7177	<u>21,7</u> 18,1
	206Pb*	207Pb*	2615,0	2629,7	2643,5	2668.4	2673,3	2683,3	2695,8	2737,0	2746,1	3096,3		41,4	21,9	33,4	18,4	8,8	15,8	2'/	11,9	0,00 85.6	48.7	377	46.3	50.0	503,0	220,9	211,5	767,0	139,1	49/	119.8	439,8	174,8	407,6	1735,7	1736,2	1749,1	1834,7	1858,1	1884,1	1888,4	1894,9	1090/2	1906,9
ges (Ma)	+1	(Ma)	14,8	33,2	13,/	21.0	19,8	16,7	13,3	33,4	13,4	13,7		10,9	56,8	76,3	78,9	70,1	23,5	31,2	22,6	15 1	012	23.6	10.7	57.5	12,6	5,5	6,5	23,8	13,5	0,10	285	33,1	27,2	23,6	19,6	17,7	13,9	19,2	28,6	23,3	26,2	16,6	0,01	13,0 14,6
Apparent a	207Pb*	235U	2591,4	2612,2	2592,0	2680.0	2675,9	2691,2	2675,3	2761,7	2732,3	3057,0		150,0	150,4	153,0	152,8	153,8	160,5	160,9	214,4	71020	220,0	220,2	220,02	234.6	282.7	253,4	257,2	325,3	255,0	C,062 N C 7 C	322.0	375,6	344,7	439,0	1737,8	1739,8	1741,5	1673,9	1873,2	1836,3	1781,9	1914,2	1,0001	1916,1
	+1	(Ma)	21,1	34,9	20,9	37.7	40,6	29,8	21,7	76,6	22,1	24,0		3,0	6,3	14,7	8,5	6'9	5,1	/,3	4,8	0,0	69	43	46	0, 8,8	5,5	4,1	3,0	15,4	5,4 0,01	e 0 ا 0,9	40	11,9	9,7	9,0	17,2	26,5	17,4	28,6	44,7	37,9	37,7	27,5	172	22,6
	206Pb*	238U	2561,4	2589,6	2526,/	2695.3	2679,4	2701,6	2648,3	2795,6	2713,7	2997,5		156,9	158,7	160,8	161,6	163,4	170,5	1/1,9	233,3	241,0	246.1	2481	748.8	253.4	256,8	256,9	262,2	267,0	///07	0,020	350.6	365,3	370,4	445,0	1739,6	1742,8	1735,1	1548,6	1886,8	1794,4	1692,4	1932,2	1011 6	1911,0 1924,5
	error	corr.	0,63	0,46	0,68	0.77	0,88	0,76	0,71	0,96	0,71	0,71		0,25	0,1	0,17	0,1	0,09	0,19	0,2	0,18	0,19	CC'0	0 15	0.0	0.13	0,43	0,66	0,41	0,69	0,34	0,24	0 11	0,32	0,29	0,31	0,48	0,81	0,69	0,89	0,82	0,88	0,81	0,85	27/0	εα'υ 0,8
	+1	(%)	-	1,6		1.7	1,9	1,3	-	3,4	-	-		1,9	4	9,3	5,3	4,3	m (4,3	2,1	0,1	0,2	18	ο τ	3.5	2,2	1,6	1,2	5,9	7,1	0,4 1 0	1 2	3,4	2,7	2,1	1,1	1,7	1,1	2,1	2,7	2,4	2,5	1,6	<u>,</u> .	1,4 -
ratios	206Pb*	238U	0,48785	0,49438	0,4/98/	0.51909	0,51534	0,52057	0,50805	0,54291	0,52342	0,592		0,02465	0,02492	0,02527	0,02539	0,02567	0,02679	0,02/02	0,03685	0.03874	0.03892	0.03974	0.03936	0.04009	0,04064	0,04065	0,04151	0,04228	0,0424	005300	0.05589	0,0583	0,05915	0,07147	0,30977	0,31043	0,30885	0,27153	0,34003	0,32096	0,30022	0,34949	0,34717	0,3479
Isotopic	+1	(%)	1,6	3,5	- 2, L	2.2	2,1	1,8	1,4	3,5	1,4	1,4		7,8	40,6	53,6	55,5	49	15,8	20,9	11,7	C,0 5 7	13.3	11 5	0,2	27.4	5,1	2,5	2,9	8,6	9 7	2 7	10.3	10,6	9,3	6,6	2,4	2,1	1,7	2,3	3,3	2,7	3,1	1,9	0,1	1,7
	207Pb*	235U	11,83485	12,09953	11,84248	13.00456	12,94821	13,15951	12,94075	14,17815	13,74526	19,30147		0,15919	0,15966	0,16265	0,16241	0,16358	0,17123	0,1/1/4	0,23515	0,24029	0.752.09	0.25255	0.25253	0.2599	0,32104	0,28342	0,28822	0,37761	0,28542	0,133910	037315	0,44758	0,40424	0,54084	4,53723	4,54824	4,55706	4,19924	5,32685	5,10123	4,78307	5,58793	01100,0	5,5997
	U/Th		2,7	4,6	2,1	1.6	1,2	0,9	1,5	1,4	1,6	3,6		-	1,4	1,3	1,4	1,2	1,7	2,6	1,4	2,7	1.7	07	0.0	1.2	2,6	1,2	0,5	2,3			0.8	0,6	1,1	1,9	1,3	1,7	2,4	19,7	0,8	5,8	3,4	1,6	7'0	1,7
	206Pb	204Pb	97124	85572	83/25	12645	110030	29013	91953	50276	42597	93698	a	4702	943	960	555	1529	2360	1/38	235/	0/CC 7017	7417	3580	4721	3003	2163	8892	8609	1515	38/2	2/15	3671	2223	2904	5880	9040	9573	12385	56032	32453	66914	8139	36174	21102	42128 16952
	U (ppm)		335	246	325	52	298	90	298	129	116	317	7 Chukotka	192	58	38	30	99	62	56	6/	137	22	85	176	75	234	331	176	574	142	112	89	41	51	62	30	31	47	858	144	224	24	115	70	7 5
	Analysis		GB9986-11	GB9986-35	GB9986-19 CD0006 F1	GB9986-52	GB9986-104	GB9986-40	GB9986-15	GB9986-66	GB9986-68	GB9986-45	Sample CH04 ELM	CH04ELM7-4	CH04ELM7-40	CH04ELM7-61	CH04ELM7-55	CH04ELM7-49	CH04ELM7-26	CH04ELM/-12	CH04ELM7-35	CH04ELINI/-40	CH04FI M7-44	CH04FI M7-109	CH04FI M7-20	CH04ELM7-20	CH04ELM7-84	CH04ELM7-16	CH04ELM7-36	CH04ELM7-9	CH04ELM/-83		CH04FI M7-73	CH04ELM7-81	CH04ELM7-86	CH04ELM7-67	CH04ELM7-38	CH04ELM7-92	CH04ELM7-98	CH04ELM7-94	CH04ELM7-31	CH04ELM7-33	CH04ELM7-1	CH04ELM7-65		CH04ELM1/-85 CH04ELM7-78

	Age of youngest cluster (Ma)	(3+ analyses)																																												
	+1	(Ma)	18,9	18,0	39,1	35,2	21.5	22,6	34,8	32,5	20,7	17,9	24,4	23,5	18.0	21.2	25,4	42,5	23,5	17,9	40,1	32,2	41,0	17,9	17,9	28,3	33,5	25,3	44,9	347	23.4	44,7	18,0	30,7	47,0	18,6	32,0	26,9	34,4 33 8	17.8	38,3	41,8	18,6	25,2	32,2	25,2
	Best age	(Ma)	1910,2	1910,3	1911,0	1911,6	1915.3	1915,4	1916,7	1918,3	1918,6	1919,2	1921,0	1924,7	1977 2	1928.6	1928,6	1928,9	1931,1	1933,7	1935,1	1935,2	1935,7	1936,4	1938,5	1939,4	1942,6	1943,4	1944,1	1950,1	1951.8	1953,3	1955,0	1956,7	1959,5	1959,9	1960,6	1962,8	1967 3	1976.7	1977,2	1982,1	1999,8	2000,0	2004,3	2004,4 2004,6
	+1	(Ma)	18,9	18,0	39,1	35,2	21.5	22,6	34,8	32,5	20,7	17,9	24,4	23,5	180	21.2	25,4	42,5	23,5	17,9	40,1	32,2	41,0	17,9	17,9	28,3	33,5	25,3	44,9	347	23.4	44,7	18,0	30,7	47,0	18,6	32,0	26,9	34,4 33.8	17.8	38,3	41,8	18,6	25,2	32,2	25,2
	206Pb*	207Pb*	1910,2	1910,3	1911,0	1911,6	1915,3	1915,4	1916,7	1918,3	1918,6	1919,2	1921,0	1924,7	1977 2	1928.6	1928,6	1928,9	1931,1	1933,7	1935,1	1935,2	1935,7	1936,4	1938,5	1939,4	1942,6	1943,4	1944,1	1950,1	1951.8	1953,3	1955,0	1956,7	1959,5	1959,9	1960,6	1962,8	19673	1976.7	1977,2	1982,1	1999,8	2000,0	2004,3	2004,6
jes (Ma)	+I	(Ma)	15,3	13,5	21,1	38,6	16.2	14,4	18,8	19,7	17,0	12,1	16,6	14,2	15.6	13.3	22,0	22,8	16,1	12,2	22,5	18,7	22,8	12,7	12,2	21,4	20,2	15,2	2///	38,3 18 0	15.3	27,8	12,5	18,8	25,3	12,5	57,4	21,6	22,5	12.2	36,2	22,8	14,1	18,5	18,1	16,5
Apparent ag	207Pb*	235U	1917,3	1931,8	1918,5	1878,0 1893 4	1936.4	1899,4	1920,0	1921,8	1921,1	1890,8	1908,1	1919,0	1975 9	1909.4	1929,2	1928,1	1916,9	1924,0	1924,8	1919,8	1945,9	1939,6	1932,0	1927,3	1933,8	1912,4	1926,9	1899,5	1909.6	1917,5	1942,6	1985,8	1941,1	1945,7	1920,3	1955,0	19/2/0	1944.6	1971,1	1996,3	2032,4	2034,5	1979,2	2018,7
-	+1	(Ma)	23,8	20,1	18,6	65,3 16.3	24.2	18,1	16,6	23,1	26,5	16,2	22,5	16,6 75.4	749	16,4	35,3	19,5	21,9	16,6	21,9	19,8	22,0	18,0	16,6	31,7	23,2	17,4	33,0	167	19.4	33,3	17,2	22,6	21,3	16,7	104,6	33,2	37.4	16.6	60,5	19,5	21,5	27,4	17,0	21,4
	206Pb*	238U	1923,9	1951,9	1925,5	1847,8	1956.2	1884,8	1923,1	1925,0	1923,3	1865,0	1896,3	1913,7	1924 6	1891.7	1929,7	1927,3	1903,7	1915,0	1915,2	1905,6	1955,4	1942,6	1925,9	1916,1	1925,6	1883,9	1910,9	1853,6	1870.9	1884,6	1931,0	2013,8	1923,9	1932,3	1883,2	1947,7	1980,1	1914.6	1965,2	2010,0	2064,7	2068,7	1955,4	2032,7
	error	corr.	0,8	0,77	0,46	0,9	0.77	0,66	0,46	0,61	0,81	0,71	0,71	0,61	0,83	0.65	0,83	0,44	0,71	0,71	0,51	0,56	0,5	0,73	0,71	0,77	0,6	0,6	0,62	0,69	0.67	0,63	0,72	0,61	0,44	0,69	0,96	0,79	00/0	0.71	0,86	0,43	0,76	0,74	0,49	0,/6
	+1	(%)	1,4	1,2	1,1	4,1	1,4	1,1	-	1,4	1,6	-	1,4	- u	, r	-	2,1	1,2	1,3	-	1,3	1,2	1,3	1,1	-	1,9	1,4	1,1	7	3,1	1.2	2	1	1,3	1,3	-	6,4	1 2)')))	1	3,6	1,1	1,2	1,5		1,2
ratios	206Pb*	238U	0,34777	0,35364	0,34809	0,33194	0.35452	0,33961	0,34759	0,348	0,34764	0,3355	0,34201	0,34562	0.34791	0.34105	0,34898	0,34846	0,34354	0,34591	0,34595	0,34394	0,35437	0,35167	0,34818	0,34613	0,34812	0,33943	0,34504	0,33313	0.33673	0,33957	0,34925	0,36669	0,34777	0,34953	0,33927	0,35274	16866,0	0.34582	0,35643	0,36589	0,37753	0,37838	0,35436	0,3/0/1
Isotopic	+1	(%)	1,8	1,6	2,4	4,5	1.9	1,7	2,2	2,3	2	1,4	1,9	1,6	- 1	1.5	2,6	2,6	1,9	1,4	2,6	2,2	2,6	1,5	1,4	2,5	2,3	1,8	3,2	4,5	1.8	3,2	1,4	2,2	2,9	1,4	6,7	2,5	0'7	1,4	4,2	2,6	1,6	2,1	2,1	1,9
	207Pb*	235U	5,60795	5,7031	5,61568	5,35697	5.73323	5,49239	5,62561	5,63715	5,63235	5,43746	5,54845	5,61864	5 66374	5.55654	5,68546	5,67826	5,6049	5,65166	5,65659	5,62416	5,7965	5,75441	5,70411	5,67337	5,71605	5,57591	5,6/028	5,493 5,7646	5.55766	5,6093	5,77452	6,06877	5,76451	5,79516	5,62719	5,85782	5 83560	5.78789	5,96717	6,14226	6,40132	6,41633	6,02352	6,30195 6,10672
	U/Th		1,5	2	1,8	1,1 2.5	3.2	5,4	1,3	1,5	1,7	1,4	0,9	0,5	0,0	0.6	4,9	0,9	5,5	3,1	1,6	4,5	7,4	0,7	2,5	5,1	3,8	1,1	- - -	2,7	1.2	2,8	1,3	2,5	2	4,3	1,2	1,5	- '- - '-	<u>,</u> 8,	4	1,8	1,8	3,4	2,7	2,8
	206Pb	204Pb	20781	60323	55857	38242	79596	12670	58246	43479	31543	78465	33424	34727	37379	36892	114482	82289	124343	53762	35858	63509	119961	64463	75510	157002	90528	33221	1904/6	8358	32794	22678	29817	14936	31482	103048	40481	6515	41037	46781	58498	115927	28855	133380	30965	37114
	U (ppm)		51	159	181	141	294	34	155	66	129	286	86	165	95	104	510	266	408	146	138	262	290	209	326	475	505	132	825	42	159	301	78	48	90	487	168	22	101	239	326	467	87	242	121	89
	Analysis		CH04ELM7-91	CH04ELM7-34	CH04ELM7-52	CH04ELM7-7 CH04ELM7-28	CH04ELM7-80	CH04ELM7-93	CH04ELM7-100	CH04ELM7-110	CH04ELM7-64	CH04ELM7-24	CH04ELM7-107	CH04ELM7-17	CH04FI M7-46	CH04ELM7-30	CH04ELM7-70	CH04ELM7-95	CH04ELM7-53	CH04ELM7-2	CH04ELM7-99	CH04ELM7-18	CH04ELM7-79	CH04ELM7-88	CH04ELM7-19	CH04ELM7-6	CH04ELM7-74	CH04ELM7-14	CH04ELM/-9/	CH04ELM7-10	CH04ELM7-58	CH04ELM7-108	CH04ELM7-27	CH04ELM7-57	CH04ELM7-60	CH04ELM7-25	CH04ELM7-15	CH04ELM7-63	CH04ELM/-51 CH04ELM7-13	CH04ELM7-37	CH04ELM7-22	CH04ELM7-50	CH04ELM7-42	CH04ELM7-82	CH04ELM7-23	CH04ELM/-45 CH04ELM7-105

	Age of youngest cluster (Ma)	(3+ analyses)																		159 (8); 175 (8)																											
	+1	(Ma)	37,5	24,9	18,3	40,4 22,6	75,2	34,4	35,9	27,4	36,6	35,9	30,2	23,6	C,01	10,4	7/01	7'N7	20,5		8,9	3,0	11,6	8,6	12,6	9,7	9,3	7,1	4,8	2,5	0,3 5,3	C'2	10,9	4,2	5,8	7,4	2,9	4,3	10,9	2,7	8,7	15,9	9,8	3,4	8,8	C'7	25,8
-	sest age	(Ma)	2005,1	2006,9	2018,5	2032,7	2058,8	2075,7	2125,9	2220,6	2381,9	2408,1	2720,6	2723,0	2/30,0	1,2012	1,4012	27001	7/90,1		156,1	158,6	159,1	163,6	166,7	169,2	169,3	173,3	174,5	176,5	232,0	242.2	244,1	244,5	244,6	246,0	248,2	251,3	253,0	257,1	257,5	266,6	270,8	307,3	345,5	122,0	1864,9
-	+	(Ma)	37,5	24,9	18,3	40,4 22,6	75,2	34,4	35,9	27,4	36,6	35,9	30,2	23,6	C,01	10,4	7/01	707	20,5		855,4	427,1	1176,4	636,7	423,8	1383,4	314,8	202,4	136,4	202,1	6/2CI	295.8	297,2	95,9	416,5	229,7	55,4	128,0	438,5	93,4	50,5	411,9	268,7	108,1	121,7	0,001	25,8
	206Pb*	207Pb*	2005,1	2006,9	2018,5	2032,7	2058,8	2075,7	2125,9	2220,6	2381,9	2408,1	2720,6	2723,0	2/30,0	1,2012	1,40,2	0,417	1/06/2		0,6	50,2	58,7	20,9	48,8	36,1	26,9	156,6	46,6	43,0	216.8	54.9	15,4	157,1	52,9	29,9	237,8	240,6	42,5	170,1	311,1	303,0	233,1	200,8	233,4	200,4	1864,9
jes (Ma)	+	(Ma)	22,8	20,9	17,7	26,4 14.4	73,3	34,0	33,6	27,9	26,7	23,8	19,7	17,2	15,3	0,01	17.0	0,11	24,5		47,9	25,2	67,4	38,3	28,3	80,6	21,0	15,2	9,7	13,2	10.2	25.7	26,4	9,4	35,7	20,4	5,9	12,9	39,1	9,1	9,5	45,0	28,6	12,3	16,6	10,0	41,U 16,3
pparent ag	207Pb*	235U	2004,8	1982,6	2022,7	1998,3 2045_0	1861,8	2040,1	2079,6	2095,5	2373,6	2452,6	2682,8	2683,4	2093,8	1,0012	4'/C/Z	2/04,0	2//4,0		146,9	152,0	152,9	154,7	159,2	160,7	160,2	172,1	166,0	167,6	7395	225.5	223,8	236,4	227,4	226,6	247,2	250,2	233,6	248,7	262,9	270,3	266,9	295,3	331,4	341,8	1875,4
A	+1	(Ma)	26,5	32,9	30,3	33,7	110,7	57,4	55,6	46,2	38,8	30,7	21,6	24,6	1/17	24,0 7,15	C'77	7/70	47,2		8,9	3,0	11,6	8,6	12,6	9,7	9,3	7,1	4,8	2,5	0,0 5,3	C/2	10,9	4,2	5,8	7,4	2,9	4,3	10,9	2,7	8,7	15,9	9,8	3,4	8,8	C'7	40,4 20,5
	206Pb*	238U	2004,5	1959,4	2026,8	2032.5	1690,6	2005,0	2033,1	1970,4	2363,9	2506,6	2632,9	2631,0	2040,U	1,0012	2/01/2	0,6612	5,55/2		156,1	158,6	159,1	163,6	166,7	169,2	169,3	173,3	174,5	176,5	232,0 741 8	242.2	244,1	244,5	244,6	246,0	248,2	251,3	253,0	257,1	257,5	266,6	270,8	307,3	345,5	1203,0	1/32,/ 1884,9
-	error	corr.	0,59	0,81	0,86	0,66	0,87	0,86	0,84	0,86	0,67	0,57	0,48	0,62	0,/9	1/10	10,0	0//0	0,/4		0,16	0,11	0,16	0,2	0,4	0,11	0,39	0,43	0,44	0,17	0,39	0.24	0,35	0,39	0,14	0,31	0,45	0,3	0,23	0,26	0,84	0,32	0,3	0,24	0,44	0,04	0,66
-	+	(%)	1,5	1,9	1,7	1 2	7,4	3,3	3,2	2,7	2	1,5	-	-, , , ,	λ 1 - 1				יע		5,7	1,9	7,4	5,3	7,6	5,8	5,5	4,2	2,8	1,4	2,8	1 1 1	4,6	1,7	2,4	3,1	1,2	1,8	4,4	1,1	3,4	6,1	3,7	1,1	2,6	3,0	3,1 1,3
atios	206Pb*	238U	0,36472	0,3552	0,36945	0,3564	0,29985	0,36482	0,37079	0,35752	0,44296	0,47526	0,50445	0,50402	77105/0	0/170/0	0,00400	100000	7075C/N		0,02451	0,02491	0,02498	0,0257	0,02619	0,0266	0,02662	0,02724	0,02744	0,02776	0,03873	0.03829	0,0386	0,03866	0,03867	0,0389	0,03926	0,03975	0,04002	0,04069	0,04076	0,04222	0,0429	0,04883	0,05505	82050,0	0,33964
sotopic r	+	(%)	2,6	2,4	7	163	8,6	3,9	3,8	3,1	2,9	2,6	2,1	1,8	0, 1	0 	C, L	א <u>ר</u> אר	7,0		35	17,9	47,4	26,7	19,2	54,1	14,2	9'6	6,3	8,6	48	12.7	13,1	4,4	17,5	10,1	2,7	5,8	18,7	4,1	4,1	18,9	12,2	4,8	5,9	/'C	4, ⁴ 1,9
	207Pb*	235U	6,20238	6,04669	6,33077	6,15631 6,49308	5,25637	6,45715	6,7529	6,87514	9,35644	10,19447	13,04338	13,05109	13,1905/	10//01	14,1134	30707011	14,5720		0,15564	0,16148	0,16255	0,16456	0,1697	0,17144	0,17088	0,18476	0,17764	0,17944	0,24328	0.24872	0,24661	0,26221	0,25099	0,25007	0,27568	0,27947	0,25865	0,27753	0,29552	0,30505	0,3007	0,33749	0,38589	0,40017	4,6129 5,34097
	U/Th		1,4	1,3	1,2	1,3	3,7	2,2	0,7	13,4	4,5	1,8	1,4	1,5	7,1	0 C	C, L		-,-		0,7	0,8	0,9	1,1	1,8	1,4	1,4	1,2	3,3	1,1	1,8	0,80	0,5	1,8	0,6	0,9	1,5	1,5	0,7	0,5	0,7	3,3	0,8	1,5	1,2	0,8	0,/ 2,9
	206Pb	204Pb	38884	29344	37630	163931 58804	8718	41358	46084	211176	124138	77484	94639	39136	52030	10070	121261	1000001	600001		1521	3062	769	1143	1206	872	1732	1976	3997	4021	2910	2310	2008	5073	1786	2769	7711	4209	1307	6665	8985	624	1922	7215	8494	3537	69820
	U (ppm)		79	96	89	517	647	155	338	761	445	359	357	69	115		140	407	677	ukotka	57	157	36	55	47	40	60	111	155	163	101	85	41	134	64	95	432	210	42	218	597	83	79	188	148	0/	194
	Analysis		CH04ELM7-106	CH04ELM7-68	CH04ELM7-103	CH04ELM7-101 CH04FI M7-72	CH04ELM7-66	CH04ELM7-11	CH04ELM7-104	CH04ELM7-90	CH04ELM7-3	CH04ELM7-39	CH04ELM7-56	CH04ELM7-32	CH04ELM/-89					Sample 04JT 54 Ch	04JT54C-6	04JT54C-98	04JT54C-45	04JT54C-40	04JT54C-97	04JT54C-13	04JT54C-7	04JT54C-32	04JT54C-101	04JT54C-11	04J154C-49 04IT54C-25	04JT54C-105	04JT54C-5	04JT54C-102	04JT54C-68	04JT54C-44	04JT54C-100	04JT54C-67	04JT54C-10	04JT54C-46	04JT54C-56	04JT54C-43	04JT54C-92	04JT54C-47	04JT54C-8	04J154C-15	04J154C-21 04JT54C-36

	Age of youngest cluster (Ma)	(3+ analyses)																																												
	+1	(Ma)	39,9	26,9	36,0	18,1	22,3	27,5	18,0	19,0	21,0	17,9	17,9	25,3	33.0	25,6	33,7	17,9	20,6	18,0	20,6	19,9	20,4	28,3	18,6	26,1	28,8	29,0	17,9	39,7	19.1	18,3	17,9	28,6	41,8	17,9	22,0	17,9	6/1	33.9	26.1	33.7	27,6	28,1	25,1	17,8 26,8
-	Best age	(Ma)	1868,6	1882,5	1885,3	1885,7 1899.5	1906,7	1907,0	1910,2	1910,9	1913,7	1913,9	1917,3	1917,3	19197	1920,5	1922,5	1923,1	1924,7	1925,8	1927,1	1928,2	1929,1	1929,7	1936,4	1937,1	1937,5	1937,8	1939,7	1941,4	1944.7	1945,9	1948,8	1949,9	1950,9	1951,5	1951,6	1952,5	1054.0	1955.4	1956.5	1960.1	1977,7	1990,0	1990,6	1992,9
-	+1	(Ma)	39,9	26,9	36,0	18,1	22,3	27,5	18,0	19,0	21,0	17,9	17,9	25,3	33.0	25,6	33,7	17,9	20,6	18,0	20,6	19,9	20,4	28,3	18,6	26,1	28,8	29,0	17,9	39,7	19.1	18,3	17,9	28,6	41,8	17,9	22,0	17,9	6/1	33.9	26.1	33.7	27,6	28,1	25,1	17,8 26,8
-	206Pb*	207Pb*	1868,6	1882,5	1885,3	1885,7 1899.5	1906,7	1907,0	1910,2	1910,9	1913,7	1913,9	1917,3	1917,3	19197	1920,5	1922,5	1923,1	1924,7	1925,8	1927,1	1928,2	1929,1	1929,7	1936,4	1937,1	1937,5	1937,8	1939,7	1941,4	1944.7	1945,9	1948,8	1949,9	1950,9	1951,5	1951,6	1952,5	1953,2	1955.4	1956.5	1960.1	1977,7	1990,0	1990,6	1992,9 1999,4
ges (Ma)	+1	(Ma)	27,2	17,4	27,3	14,3	15,5	17,5	12,2	12,6	13,2	12,2	22,1	15,3	34.0	18,1	40,2	13,1	13,9	12,1	13,2	32,9	13,3	17,9	13,8	15,5	16,3	19,6	18,5	24,5	12.7	14,2	12,2	16,3	24,2	12,3	28,5	29,1	1 1 1	37.9	17.1	19.6	19,5	19,7	24,2	22,5
Apparent a	207Pb*	235U	1883,2	1890,2	1912,0	1932,5	1899,5	1917,4	1900,8	1915,5	1899,2	1924,4	1924,0	1963,3	1908 1	1915,6	1797,9	1917,5	1920,4	1856,7	1931,1	1920,5	1936,7	1933,4	1982,6	1956,6	1921,5	1947,1	1966,7	1905,5 1740 E	1940,8	1958,2	1935,3	1930,3	1920,9	1963,7	1889,0	1904,9	1 924,8	1914.0	1946.7	1914.5	1959,0	2007,0	1956,1	1941,0 1984,7
	+1	(Ma)	37,5	22,5	41,0	22,5	21,5	22,2	16,4	16,6	16,4	16,7	39,3	18,2	574	25,5	65,3	19,0	18,7	15,7	16,7	60,4	17,3	22,4	20,8	17,5	16,5	26,7	32,2	28,9	16.7	21,6	16,6	16,5	25,4	17,0	49,3	52,3	16,4	64.7	222	20.4	27,3	27,8	40,0	39,5
-	206Pb*	238U	1896,4	1897,1	1936,7	1976,4	1893,0	1926,9	1892,2	1919,7	1886,0	1934,2	1930,2	2007,2 1005 E	1897.6	1911,1	1692,4	1912,4	1916,5	1795,7	1934,7	1913,4	1943,7	1936,9	2027,3	1975,2	1906,8	1955,8	1992,5	1872,7	1937.1	1969,8	1922,6	1912,0	1893,3	1975,2	1832,5	1861,4	1000.2	1875.9	1937.4	1872.8	1941,3	2023,6	1923,6	1892,8 1970,6
-	error	corr.	0,72	0,67	0,77	0,8	0,73	0,66	0,71	0,69	0,65	0,71	0,92	0,6	0.88	0,73	0,92	0,75	0,7	0,71	0,66	0,96	0,67	0,65	0,75	0,58	0,53	0,7	0,88	0,63	0.68	0,78	0,71	0,53	0,55	0,71	0,93	0,96	0,/1	0.0	0.67	0.55	0,72	0,71	0,86	0,92
	+1	(%)	2,3	1,4	2,5	1,3	1,3	1,3	-	-	-	-	2,4	1,1	- 5 5	1,5	4,4	1,1	1,1		-	3,6	-	1,3	1,2	- '	-	1,6	1,9	1,8		1,3	-		1,5	-	3,1	3,2		- 4	1.3	1.0	1,6	1,6	2,4	2,5
ratios	206Pb*	238U	0,34202	0,34217	0,35044	0,35879	0,34132	0,3484	0,34115	0,34689	0,33987	0,34991	0,34908	0,3653	0 34776	0,34508	0,30022	0,34535	0,34622	0,32122	0,35003	0,34557	0,35192	0,35048	0,36955	0,35852	0,34418	0,35445	0,36219	0,33709	0.35053	0,3574	0,34749	0,34528	0,34137	0,35854	0,32879	0,33477	0,34245	0,33777	0.35059	0.33711	0,35141	0,36877	0,3477	0,34127 0,35756
lsotopic	+1	(%)	3,2	2	3,2	1,7	1,8	2	1,4	1,5	1,5	1,4	2,6	1,8	- °	2,1	4,8	1,5	1,6	1,4	1,5	3,8	1,5	2,1	1,6	1,8	1,9	2,3	2,1	2,8	1.5	1,6	1,4	1,9	2,8	1,4	3,3	3,4	- r 4 v	44	C	23	2,2	2,2	2,8	2,6
	207Pb*	235U	5,38941	5,43356	5,5733	5,70733	5,49328	5,60823	5,50151	5,596	5,49126	5,65437	5,65147	5,91397 5 50101	5 54847	5,59686	4,87467	5,60915	5,62827	5,22522	5,69799	5,62885	5,73505	5,71353	6,04706	5,86891	5,63543	5,80444	5,93748	5,53125	5.76243	5,87934	5,72581	5,69273	5,63135	5,91648	5,42609	5,5276	2,00003	5,58615	5.8017	5,58992	5,88496	6,21835	5,86506	5,76398 6,06145
	U/Th		7,1	1,7	0,7	0,4	0,7	0,7	1,3	1,3	3,6	0,7	5,2	2,7	t,0 0	3,3	0,4	0,6	11,2	1,6	1,3	3,4	1,1	0,3	6	7,1	309	2,6	2,9	10,9	4,0	3,7	6,1	3,6	0,8	1,8	19,7	m,		, r	10.1	2.3	2,5	2,1	4	3,9 29,8
	206Pb	204Pb	4036	9011	33547	19535	36848	13322	24085	54577	59316	92351	101741	44550	46860	80428	15719	35094	42141	33205	42482	96175	60260	40116	190411	183934	172165	38281	44067	180781	21197	49694	196821	117509	49255	40118	154385	50959	22320	45098	61312	78118	7878	107019	41331	44320 83562
	U (ppm)		13	27	152	55	120	50	89	211	253	400	352	215	2001	302	60	240	149	207	205	2014	182	178	1043	694	659	210	166	714	47	635	1240	440	213	143	651	839	130	138	387	744	121	389	98	763 654
	Analysis		04JT54C-17	04JT54C-81	04JT54C-51	04JT54C-74 04IT54C-91	04JT54C-16	04JT54C-107	04JT54C-66	04JT54C-53	04JT54C-61	04JT54C-69	04JT54C-27	04JT54C-78	04JT54C-03	04JT54C-59	04JT54C-58	04JT54C-75	04JT54C-95	04JT54C-29	04JT54C-72	04JT54C-77	04JT54C-22	04JT54C-28	04JT54C-33	04JT54C-76	04JT54C-14	04JT54C-35	04JT54C-55	04J154C-57	04JT54C-02 04JT54C-73	04JT54C-85	04JT54C-82	04JT54C-37	04JT54C-89	04JT54C-80	04JT54C-104	04JT54C-71	04J154C-54	041T54C-86	04JT54C-110	04IT54C-87	04JT54C-50	04JT54C-30	04JT54C-20	04JT54C-83 04JT54C-108

					lsotopic	ratios					Apparent ag	ges (Ma)					
Analysis	U (ppm)	206Pb	U/Th	207Pb*	+I	206Pb*	+1	error	206Pb*	+1	207Pb*	+1	206Pb*	+1	Best age	+I	Age of youngest cluster (Ma)
		204Pb		235U	(%)	238U	(%)	corr.	238U	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(3+ analyses)
04JT54C-96	533	142244	4	6,13206	1,4	0,36073	1	0,71	1985,6	17,1	1994,8	12,3	2004,3	17,8	2004,3	17,8	
04JT54C-12	198	62275	e	6,14509	2,6	0,36129	2,2	0,86	1988,3	38,0	1996,7	22,5	2005,3	22,9	2005,3	22,9	
04JT54C-38	315	84250	13,6	6,26682	1,7	0,36838		0,58	2021,8	17,4	2013,8	15,0	2005,7	24,7	2005,7	24,7	
04JT54C-1	219	34558	0,9	6,04867	2,3 1 8	0,35478	1,7	0,72	1957,4	28,0	1982,9	20,2	2009,6	28,8	2009,6	28,8	
04JT54C-41	53	16551	- 1-	6.26338	2 1.7	0.3665	<u>,</u> –	0.65	2012.9	17.3	2013.3	13.5	2013.8	20.8	2013.8	20.8	
04JT54C-93	67	19702	1,9	6,29086	1,8	0,36626	1,4	0,76	2011,8	23,6	2017,2	15,8	2022,7	20,9	2022,7	20,9	
04JT54C-2	160	35353	2,1	6,09983	1,9	0,35388	-	0,51	1953,1	16,8	1990,2	17,0	2029,0	29,6	2029,0	29,6	
04JT54C-109	72	18343	1,5	6,08381	2,9	0,35268	2,2	0,75	1947,4	36,5	1987,9	25,3	2030,4	34,0	2030,4	34,0	
04JT54C-23	282	52292	1,5	6,54897	2,1	0,37706	1,8	0,84	2062,5	31,9	2052,5	18,9	2042,4	20,5	2042,4	20,5	
04JT54C-3	390	47477	0,7	6,08557	4,3	0,34946	3,9	0,91	1932,0	65,3	1988,2	37,3	2047,1	30,8	2047,1	30,8	
04JT54C-99	205	36270	1,4	6,52273	1,6	0,37384		0,62	2047,4	17,5	2049,0	14,3	2050,5	22,6	2050,5	22,6	
04JT54C-39	143	36171	2,7	6,31114	5,5	0,36162	4,7	0,86	1989,8	81,0	2020,0	48,1	2051,0	49,1	2051,0	49,1	
04J154C-94	61	19115	n v	6,647	2,4	0,37944	6'L	0,8	20/3//	34,0	2065,6	21,3	2057,6	25,8	2057,6	25,8	
04J154C-19	151	7/177	0,1	6,18013	/'	0,35155	7, L	1//0	1942,0	20,8	2001,6	12,21	2063,8	2'17 2'17	2063,8	21,5	
04J154C-52	142	5480	- c	0,45283	7 7	0,3065/	τ, <u>Γ</u>	0,/3	2013,2	C,CC	2039,5	8/1	2066,1	24,0	2066,1	24,0	
04J154C-20	112	1/88/2	3,1	C024/,0	2'7 E 2	0,3/495	<u>ν</u> ο Γ	α'n	7,2012	27,2	20/8/2	20,3	2103,0	24,2	2103,0	24,2	
04J134C-4	211	13733		20057/1	C,U L C	012670	7 0	0,72	2/4/2	00,4	2140,4	10.0	C,4C12	0,00	C,4C12	0,00	
041754C-00	466	40398	17	9 95676	2,1	0.4469	- 1 - 1	0,00	73815	36.8	2430.7	19.4	74722	16.9	7472 7	16.9	
04 IT54C-79	477	89091	, r ,	9 5568	- 0°	0 47141	3.4	0.91	2266.9	654	7393.0	34.8	25022	269	21212	0.90	
04 IT 54C-70	702	78991	0,4	10.35296	2,6	0.44307	5	0.74	2200,2	38.7	2466.9	24.5	25574	202	2552.4	20'2	
04JT54C-103	117	46056		12.52672	3,2	0.5054	1.9	0.6	2637.0	41,9	2644.7	30,6	2650.7	43,3	2650,7	43.3	
04JT54C-42	518	125773	1,8	12,59685	2,4	0,50469	7	0,81	2633,9	42,9	2650,0	23,0	2662,3	23,7	2662,3	23,7	
04JT54C-90	132	56379	1,7	12,90989	2,6	0,50141	1,7	0,66	2619,9	36,5	2673,1	24,2	2713,6	31,8	2713,6	31,8	
04JT54C-9	613	180815	0,7	13,74048	4,8	0,51406	4,4	0,92	2673,9	96,4	2732,0	45,1	2775,2	29,7	2775,2	29,7	
04JT54C-60	88	22019	1,1	13,68398	2,2	0,51168	1,6	0,76	2663,8	35,8	2728,1	20,5	2776,0	23,2	2776,0	23,2	
04JT54C-34	131	57740	0,8	16,90101	4,6	0,58272	4,2	0,91	2959,8	98,9	2929,2	43,8	2908,3	30,1	2908,3	30,1	
Sample ELM 03 CI	424.3A																182 (4)
ELM03CH243A-96	1024	3488	0,9	0,18989	15,48	0,02688	8,72	0,56	171	15	177	25	252	147	171	15	
ELM03CH243A-21	140	836	1,6	0,21714	25,48	0,02846	3,37	0,13	181	9	200	45	426	282	181	9	
ELM03CH243A-51	266	2572	0,9	0,20842	29,46	0,02954	6,06	0,21	188	11	192	50	248	332	188	11	
ELM03CH243A-81	126	284	1,1	0,10099	63,41	0,03119	7,44	0,12	198	15	98	57	-1987	1176	198	15	
ELM03CH243A-75	580	1370	1,2	0,24756	17,04	0,03318	5,25	0,31	210	;	225	34	376	182	210	;	
ELMU3CH243A-31	110	00/1	α [′] - Γ	C249220	07 CC	0,03345	4,38 F 00	0,18	717	א ע	120	40	5/3	607	717	ן ע	
FI M03CH243A-53	216	496	2'-	0.05396	37.67	0.03995	2,68	0,20	240	7	2/1	75	ncc E	453	240	7	
ELM03CH243A-69	228	2740	0,9	0,31699	13,91	0,04097	3,86	0,28	259	10	280	33	457	148	259	10	
ELM03CH243A-92	414	1302	1,6	0,30552	20,12	0,04155	1,58	0,08	262	4	271	47	343	227	262	4	
ELM03CH243A-70	358	2810	1,2	0,31783	18,31	0,04253	3,09	0,17	269	∞	280	44	379	203	269	∞	
ELM03CH243A-32	950	2356	6'0	0,33662	11,22	0,04293	1,9	0,17	271	5	295	28	487	122	271	5	
ELM03CH243A-//	404	1590	0,9	0,33764	1/,33	0,04334	1,23	/0/0	2/4	τ, τ	295	44	4/2	191	2/4	τ, Γ	
ELMU3CH243A-10	102	8761	<u>ν</u>	0,33919	25,45	0,04364	17'5	0,14	C/7	4 5	167	71	407	202	C/2	4 5	
FI M03CH243A-13	336	1212	2,1	11600,0	71 33	0.0447	3 15	0,14	787	- 0	CTC	- 02	187	246	200	- 0	
FI MORCH243A-4	2022	1308	- 1	0 30909	18 99	0.04516	2 10	0.14	202 285		2/2	45	178	210	202 285		
ELM03CH243A-95	110	2322	0,8	0,29086	22,58	0,04536	3,96	0,18	286	,=	259	50	24	267	286	, [
ELM03CH243A-97	152	818	0,5	0,30999	26,51	0,04594	4,63	0,17	290	13	274	62	145	306	290	13	
ELM03CH243A-29	312	1518	1,2	0,3402	17,97	0,04636	4,11	0,23	292	12	297	45	338	198	292	12	

	Age of youngest cluster (Ma)	(3+ analyses)																																													
	+1	(Ma)	6	16	12	6	6	7	16	7	17	16	י <u>י</u> ד ק	20	18	32	31	15	13	25	25	15	68	45	27	61	16	40	73	33	74	9	47	32	25	35	17	18	36	24	39	7	23	57	17	1	r 6
_	Best age	(Ma)	302	307	311 316	322	331	332	343	381	381	382	412	418	427	437	451	465	487	603	659	737	1834	1850	1859	1860	1874	1877	1888	1892	1895	1900	10/1	1909	1910	1911	1914	1919	1922	1922	9761	9761	1932	1933	1934	1047	1956
-	+1	(Ma)	242	72	353	170	194	264	329	198	150	241	CCC	315	110	262	235	189	104	188	164	108	89	45	27	61	16	40	73	33	74	9	42	32	25	35	17	18	36	24	39	7 6	22.00	7	- 1	10	10
-	206Pb*	207Pb*	35	1388	357 473	475	383	503	502	359	515	1104	558	317	418	802	180	254	530	498	555	719	1834	1850	1859	1860	1874	1877	1888	1892	1895	1900	1907	1909	1910	1911	1914	1919	1922	1922	9761	9761	1932	1933	1042	7101	1956
ges (Ma)	+1	(Ma)	48	35	82	44	49	70	89	54	47	111	71	90	37	66	70	57	38	75	1	53	64	44	29	35	36	45	71	34	75	16	16	38	29	40	54	36	43	26	40	20	35	35	20	- C	25
Apparent a	207Pb*	235U	273	472	316	342	337	354	364	378	401	503	411	403	426	500	409	431	495	582	636	733	1762	1797	1823	1866	1756	1847	1849	1881	1804	1893	1840	1909	1901	1795	1877	1882	1905	1887	1010	1918	1888	1882	1001	10101	1920
	+1	(Ma)	6	16	12	6	6	7	16	7	17	16	<u>1</u>	20	18	32	31	15	13	25	25	15	33	31	26	45	58	49	45	26	65	28	90 80	44	34	41	100	60	53	26	4	32	02 5	43	30	C/ L/	45
	206Pb*	238U	302	307	311 316	322	331	332	343	381	381	382	412	418	427	437	451	465	487	603	659	737	1703	1752	1791	1870	1658	1821	1814	1872	1725	1886	1781	1909	1893	1697	1844	1849	1888	1855	10/0	1911	1848	1830	1001	2021	1888
	error	corr.	0,15	0,59	0,12	0,18	0,16	0,09	0,15	0,11	0,31	0,18	0.18	0.17	0,4	0,29	0,33	0,2	0,28	0,24	0,26	0,2	0,28	0,37	0,49	0,8/0	0,92	0,57	0,33	0,39	0,46	0,93	0,83	0,6	0,61	0,57	0,96	0,88	0,64	0,52	90,0	18/0	0,46	(0,0 2,0,0	0,94	0,92	0,94
	+1	(%)	3,05	5,47	3,89	2,87	2,8	2,05	4,69	1,94	4,48	4,42	3.63	4.84	4,3	7,61	7,03	3,31	2,82	4,28	4,07	2,1	2,24	7 2	1,68	3,66 2.78	4	3,08	2,86	1,58	4,28	1,7	0,94 1 78	2,66	2,1	2,73	6,25	3,73	3,27	1,63	3,23	1,92	1,89	2//2	2,10	4,44	2,78
ratios	206Pb*	238U	0,04791	0,04869	0,04938	0,05127	0,05267	0,05287	0,05468	0,0608	0,06089	0,06098	0.0066	0.06705	0,06855	0,07014	0,07244	0,07474	0,07854	0,09808	0,10/5/	0,12111	0,30226	0,3123	0,32025	0,32034	0,29326	0,32637	0,32499	0,33692	0,3069	0,33993	0,2/208	0,34472	0,34135	0,30105	0,3312	0,33224	0,34034	0,33336	C0C67/0	0,34513	0,33201	0,32948	0,27233	110400	0,34018
lsotopic	+1	(%)	20,44	9,32	31,56	15,63	17,49	24,06	30,28	17,65	14,36	24,51	20.70	28.17	10,77	26,17	21,35	16,81	9,91	17,59	15,55	10,42	7,88	5,37	3,43	4,22	4,36	5,37	8,6	4,04	9,26	1,82	8,4 1 01	4,46	3,46	4,78	6,52	4,22	5,14	3,12	2,44	2,30	4,1	4,2	2,3	4,00	2,97
	207Pb*	235U	0,30861	0,59243	0,36545	0,39985	0,39422	0,4177	0,43175	0,45034	0,48374	0,64167	0.53467	0.48738	0,52126	0,63693	0,49624	0,5286	0,62803	0,773	0,8/03	1,05728	4,67207	4,86933	5,02003	5,02315	4,63488	5,1669	5,17645	5,37813	4,90793	5,44981	4,42831	5,55389	5,50381	4,85632	5,35095	5,38395	5,52525	5,41156	4,80843 7,61463	2,01402	5,41866	5,38036	2628/,6	7/100/0	5,62687
	U/Th		1,8	1,4	0,6	0,9	1,7	0,8	0,9	0,9	0,8	1,8	- 6	2	0,9	1,4	-	-	1,4	6'0	- 0	0,8	· ·	1,4	1,2	2,4	e co	4	1,1	1,5	-	2,7	8,/	0,8	2	1,7	2	2,4	1,9	1,8	4,0	4,3		0,0	10,1	- '0	1,8
	206Pb	204Pb	1706	2598	1140	2826	1194	944	1506	1368	1870	1032	1234	1414	2866	804	1608	1872	3516	2242	2926	2204	1504	2552	5074	26/6	9166	5764	1284	3520	1406	23056	10658	5812	4826	2888	12366	15976	4486	6650	3450	12380	3312	4904	20008	01001	13812
	U (ppm)		50	612	194	166	538	236	60	264	264	154	260	30	176	170	146	144	306	54	108	204	9/	118	182	322	318	164	104	118	48	970	1050	160	196	218	352	538	126	190	142	300	158	154	000	167	476
	Analysis		ELM03CH243A-11	ELM03CH243A-50	ELM03CH243A-78 FI M03CH243A-30	ELM03CH243A-61	ELM03CH243A-7	ELM03CH243A-46	ELM03CH243A-58	ELM03CH243A-55	ELM03CH243A-80	ELM03CH243A-37	FI M03CH243A-10	ELM03CH243A-90	ELM03CH243A-43	ELM03CH243A-17	ELM03CH243A-20	ELM03CH243A-64	ELM03CH243A-42	ELM03CH243A-62	ELM03CH243A-68	ELM03CH243A-91	ELM03CH243A-14	ELM03CH243A-83	ELM03CH243A-85	ELM03CH243A-59	ELM03CH243A-65	ELM03CH243A-99	ELM03CH243A-6	ELM03CH243A-74	ELM03CH243A-94	ELM03CH243A-79		ELM03CH243A-87	ELM03CH243A-25	ELM03CH243A-3	ELM03CH243A-56	ELM03CH243A-38	ELM03CH243A-57	ELM03CH243A-84	ELMU3CH243A-44	ELMU3CH243A-6/	ELM03CH243A-15	ELM03CH243A-45			ELM03CH243A-47

	Age of youngest cluster (Ma)	(3+ analyses)																																														
	+1	(Ma)	28	15	12	53	21	12	70	29	44	12	34	n	ب د ر	77	000	v 10	30	7	16	13	6	4		2,1	1,6	1,9	3,7	2,3	4,1	2,4	2,9	4,3	3,8	7'0	7,0	7'7	1'7	2,0	2.8	4.1	5.8	2,9	3,0	4,0	3,2	3,9
	Best age	(Ma)	1958	1961	1961 1964	1974	1978	1979	1983	1994	2006	2013	CI 07	2023	0000	2002	2154	2334	2364	2394	2495	2630	2688	2809		148,8	157,0	189,4	224,2	233,5	237,7	240,2	252,6	258,5	201,1	1,002	0'/97	2/0,4	C/8/7	202,0 787 8	285.4	286.4	295,1	301,4	304,1	326,3	333,5	360,7
	+1	(Ma)	28	15	13	53	21	12	70	29	44	12	39	n 0	ب ب	77	σ	<u>, r</u>	30	7	16	13	6	4		522,0	513,7	135,4	67,3	62,5	717,9	257,7	85,3	70,5	7,001	240,0	189,9	0,202	1001	0,801	171.3	22.8	231.0	133,2	56,4	64,2	47,6	32,6
	206Pb*	207Pb*	1958	1961	1961	1974	1978	1979	1983	1994	2006	2013	C102	2023	0000	6502	2154	2334	2364	2394	2495	2630	2688	2809		-87,8	-16,6	151,7	320,0	250,1	-23,9	273,9	322,8	295,4	345,0	411,8	229,0	1,000	2007	3074	156.8	319,5	207.1	255,9	190,3	371,5	339,9	381,7
jes (Ma)	+	(Ma)	60	33	30	53	30	31	68	34	50	31	- +	0 1 0	24		30	22	34	25	38	27	26	12		26,9	28,8	10,0	7,1	6,1	56,6	24,4	9,0	8,2	C, 11	20,4	761	20'07	ς,δ C 0C	20,2	17.6	4.5	25.2	15,1	6,6	8,8	6,6	5,6
pparent ac	207Pb*	235U	1838	1928	1899 1957	1932	1895	1870	1929	1954	1933	1983	1002	1007	1 0/0	1 060	1986	2248	2305	2328	2495	2532	2675	2750		135,7	146,7	186,6	232,7	235,0	215,2	243,4	259,6	262,2	209,7	280,6	203,1	C'707	7/17	2/9/2	271.9	290,1	285,4	296,3	291,3	331,9	334,3	363,6
4	+1	(Ma)	100	57	52 49	35	43	53	33	38	52	55	20	00 0	00	in ro	22	45	26	50	76	51	55	26		2,1	1,6	1,9	3,7	2,3	4,1	2,4	2,9	4,3	3,8	0'7	7,0	7'7	7'7	2,0 2,1	2,8	4.1	5,8	2,9	3,0	4,0	3,2	3,9
	206Pb*	238U	1735	1898	1842 1950	1892	1820	1774	1879	1916	1866	1954	2000	1/01	1716	1863	1878	2154	2238	2254	2495	2412	2658	2671		148,8	157,0	189,4	224,2	233,5	237,7	240,2	252,6	258,5	201,1	1,002	0'/97	2/0,4	C/8/7	202,0 282,8	285.4	286.4	295,1	301,4	304,1	326,3	333,5	360,7
	error	corr.	0,91	0,9	0,93	0,34	0,75	0,94	0,25	0,57	0,55	0,92	0,5/	0//0	100	0,04	90.0	0.98	0,37	0,96	0,89	0,86	0,91	0,91		0,07	0,05	0,17	0,49	0,35	0,06	0,09	0,30	0,48	0,30	60'0	0,12	60'0	67'0	0,12	0.14	0.83	0.20	0,17	0,38	0,40	0,43	0,61
	+1	(%)	6,65	3,48	3,27 2,92	2,13	2,7	3,45	2,03	2,27	3,23	3,29	c/'I	- 4'-	2,04	0/C	3 58	2,46	1,4	2,64	3,69	2,52	2,54	1,19		1,4	1,0	1,0	1,7	1,0	1,8	1,0	1,2	1,7	ς, l	0,1	0,0) - -	0, 0	- ⁻ -	1.0	1.5	2,0	1,0	1,0	1,3	1,0	1,1
ratios	206Pb*	238U	0,30874	0,34227	0,33072 0.35327	0,34105	0,32621	0,31668	0,33837	0,34611	0,33572	0,35412	0,3038	11/67/0	0,20044	0,33516	012020	0.39683	0,41513	0,41859	0,47258	0,45376	0,51033	0,51326		0,0234	0,0246	0,0298	0,0354	0,0369	0,0376	0,0380	0,0400	0,0409	0,0413	0,0420	0,0423	0,0433	0,0441	0,0449	0.0453	0.0454	0,0468	0,0479	0,0483	0,0519	0,0531	0,0576
lsotopic	+1	(%)	7,33	3,85	3,53 3,25	6,32	3,59	3,69	8,12	4	5,9	3,56	4,/8	70,0	C0/7	11 1	3 77	2,52	3,82	2,75	4,16	2,95	2,79	1,3		21,1	21,1	5,9	3,4	2,9	29,2	11,3	3,9	3,5	4,9	10,8	11.7	\ - - -	5,5	0,0 8,3	7.4	1.8	10,1	5,9	2,6	3,1	2,3	1,8
	207Pb*	235U	5,11253	5,67884	5,48815 5,87088	5,7006	5,46387	5,30808	5,6822	5,85089	5,71082	6,04796	CUU22,0	2,10201	4,92149 5 70447	24442	6 06841	8,15079	8,67659	8,90231	10,66818	11,10609	12,93918	14,00267		0,1430	0,1554	0,2017	0,2576	0,2604	0,2361	0,2708	0,2913	0,2946	0,3042	0,3183	20000	1025,0	0,3139	012/010/	0.3070	0,3307	0.3246	0,3388	0,3323	0,3867	0,3899	0,4305
	U/Th		2,1	7	1,8	1,9	1,7	3	1,2	2,4	0,9	, -,		1,1,1	1/11	1 6	0,1	2.5	0,9	2,3	1,8	1,3	1,6	1,4		0,6	0,7	0,7	3,0	0,6	0,9	0,8	0,6	0,4	9,4 7 1	C'7	1,1	0,0	0,9	0,9 1 ع	0.8	1.2	2.0	0,9	0,7	2,0	4,4	1,7
	206Pb	204Pb	3600	8480	9956 10000	3858	5968	13414	1316	3334	2110	10176	0067	04-4	20001	2/04	12086	23240	5054	15458	9718	6112	7918	24534		936	870	2764	4594	6484	640	1544	4190	2812	4//4	77/7	4030	2/30	0000	2007	2640	30232	2960	3498	4448	6614	11318	19882
	U (ppm)		286	492	528 456	158	322	430	114	1288	140	296	107	1010	100	200	20	1306	158	522	156	214	172	384	10	113	80	289	543	439	41	95	379	211	420	2	158	151	301	138	159	2413	104	179	187	420	476	887
	Analysis		ELM03CH243A-18	ELM03CH243A-5	ELM03CH243A-63 ELM03CH243A-49	ELM03CH243A-71	ELM03CH243A-48	ELM03CH243A-33	ELM03CH243A-8	ELM03CH243A-40	ELM03CH243A-72	ELM03CH243A-35	ELMU3CH243A-89				FI M03CH243A-66	ELM03CH243A-9	ELM03CH243A-86	ELM03CH243A-26	ELM03CH243A-27	ELM03CH243A-24	ELM03CH243A-54	ELM03CH243A-98	Sample ELM06 PV	ELM06-PV10-44	ELM06-PV10-52	ELM06-PV10-43	ELM06-PV10-32	ELM06-PV10-94	ELM06-PV10-4	ELM06-PV10-12	ELM06-PV10-10	ELM06-PV10-29	ELINIU0-PV 10-75	ELMU6-PV10-60	ELMU6-PV10-65		ELINIU0-PV 10-58	ELM06-PV10-85	EL M06-PV10-2	ELM06-PV10-23	ELM06-PV10-76	ELM06-PV10-89	ELM06-PV10-63	ELM06-PV10-28	ELM06-PV10-99	ELM06-PV10-100

	Age of youngest cluster (Ma)	(3+ analyses)																																												
	+1	(Ma)	3,9	4,7	15,9	37.8	18.3	18,3	26,5	18,2	18,2	18,4	18,8	20,1 18 1	18,1	18,1	18,1	18,1	18,1	18,1	18,1	18,1	18,0	23,7	18,0	18,1	18,0	18,0	18,0	36,0	33.4	18,2	54,5	18,0	18,0	85,6	18,8	17,9	19.7	17.9	19,0	17.9	28,6	23,1	17,8	17,8
	Best age	(Ma)	397,3	488,5	656,5	814,7 1718 1	1741.1	1747,7	1769,8	1789,8	1796,4	1836,4	1841,3	1849,2	1851,5	1855,5	1855,7	1856,9	1859,4	1859,8	1859,9	1861,0	1865,7	1868,1	1870,6	1871,2	1872,8	1872,9	1874,7	1885,2	1888.7	1899,3	1902,4	1902,4	1912,2	1914,2	1920,2	1929,7	1936.4	1943.4	1943,9	1951.9	1961,2	1966,7	1970,3	1977,3
	+	(Ma)	130,0	37,4	43,6	39,6 37,8	18.3	18,3	26,5	18,2	18,2	18,4	18,8	20,1	18,1	18,1	18,1	18,1	18,1	18,1	18,1	18,1	18,0	23,7	18,0	18,1	18,0	18,0	18,0	36,0	33.4	18,2	54,5	18,0	18,0	85,6	18,8	17,9	19.7	17.9	19,0	17,9	28,6	23,1	17,8	17,8
	206Pb*	207Pb*	445,8	479,1	762,1	782,0	1741.1	1747,7	1769,8	1789,8	1796,4	1836,4	1841,3	1849,2	1851.5	1855,5	1855,7	1856,9	1859,4	1859,8	1859,9	1861,0	1865,7	1868,1	1870,6	1871,2	1872,8	1872,9	1874,7	1 885,2	1888.7	1899,3	1902,4	1902,4	1912,2	1914,2	1920,2	1929,7	1936.4	1943.4	1943,9	1951,9	1961,2	1966,7	1970,3	1977,3
ges (Ma)	+1	(Ma)	19,8	7,6	16,2	11,9	12.6	11,8	14,8	11,6	16,7	11,9	12,1	12,7	12.0	12,1	16,8	18,9	12,2	15,6	14,5	17,0	12,1	13,9	12,1	12,6	12,1	17,9	12,0	19,3	18.2	12,7	26,5	17,0	12,2	42,7	12,5	12,3	12,21	12.2	24,5	12.3	16,6	14,2	39,1	26,3 12,4
Apparent a	207Pb*	235U	404,5	486,8	680,8	806,0 1740 9	1698,8	1740,6	1789,7	1655,6	1493,1	1764,5	1778,5	1835,1 1760 8	1811.4	1857,3	1755,9	1796,7	1890,6	1606,1	1865,0	1843,4	1866,8	1815,1	1865,0	1806,8	1873,5	1608,2	1835,4	1927,9	1919.4	1914,0	1722,1	1865,7	1922,6	1869,8	1801,8	1959,8	1945,/	1912.3	1969,2	1962.2	1896,6	1947,8	1673,0	1784,3 2028,7
_	+1	(Ma)	3,9	4,7	15,9	15.4	17.0	15,2	15,8	13,8	22,1	15,0	15,1	15,9	15,5	16,1	25,6	30,7	16,6	21,0	22,2	27,5	16,2	15,5	16,2	16,9	16,3	25,1	15,7	17,0	16,8	17,8	14,0	27,4	16,7	23,6	15,6	17,1	16.0	16,3	44,8	17,0	17,4	16,7	60,4	42,5
	206Pb*	238U	397,3	488,5	656,5	1760.0	1664.7	1734,7	1806,8	1551,9	1288,8	1704,5	1725,5	1822,6	1776,8	1858,9	1673,2	1745,3	1919,2	1419,8	1869,6	1827,9	1867,7	1769,3	1860,0	1751,4	1874,1	1413,9	1801,0	1968,0	1947.9	1927,5	1577,6	1832,9	1932,2	1830,1	1701,3	1988,4	1954 0	1883.7	1993,4	1972.0	1838,2	1930,2	1446,4	1623,9 2064,8
	error	corr.	0,17	0,51	0,78	0,47	0.76	0,71	0,57	0,71	0,88	0,70	0,69	0,67	0.71	0,71	0,87	06'0	0,71	0,86	0,81	0,87	0,71	0,61	0,71	0,74	0,71	0,89	0,71	0,45	0,47	0,73	0,31	0,86	0,71	0,30	0,70	0,71	0.68	0.71	0,93	0.71	0,56	0,61	0,98	0,95
	+1	(%)	1,0	1,0	2,5	1,0	2,1	1,0	1,0	1,0	1,9	1,0	1,0	1,0	1,0	1,0	1,7	2,0	1,0	1,7	1,4	1,7	1,0	1,0	1,0	1,1	1,0	2,0	1,0	1,0	0,1	1,1	1,0	1,7	1,0	1,5	1,0	1,0	0, 0	1,0	2,6	1,0	1,1	1,0	4,7	3,0
ratios	206Pb*	238U	0,0636	0,0787	0,1072	0,1347	0.2946	0,3088	0,3235	0,2722	0,2213	0,3027	0,3069	0,3268	0.3173	0,3342	0,2964	0,3109	0,3468	0,2464	0,3365	0,3278	0,3361	0,3158	0,3345	0,3122	0,3374	0,2452	0,3223	0,3570	0.3528	0,3485	0,2773	0,3289	0,3495	0,3283	0,3020	0,3613	0 3541	0.3394	0,3624	0,3579	0,3300	0,3491	0,2515	0,2865 0,3775
lsotopic	+1	(%)	5,9	2,0	3,3	2,1	1.5	1,4	1,8	1,4	2,1	1,4	1,4	1,5	1,4	1,4	2,0	2,2	1,4	1,9	1,7	2,0	1,4	1,6	1,4	1,5	1,4	2,2	1,4	2,2	2.1	1,5	3,2	2,0	1,4	5,0	1,5	1,4	- 1 4 1	5, L	2,8	1,4	1,9	1,6	4,8	3,1 1,4
	207Pb*	235U	0,4895	0,6152	0,9552	1,2117	4.3284	4,5526	4,8273	4,1064	3,3511	4,6849	4,7637	5,0939	4,9533	5,2286	4,6366	4,8678	5,4363	3,8636	5,2763	5,1442	5,2869	4,9752	5,2762	4,9262	5,3288	3,8737	5,0959	5,6775	5,6213	5,5862	4,4519	5,2803	5,6422	5,3058	4,8975	5,8901	7038 5 7038	5.5751	5,9546	5,9067	5,4747	5,8096	4,1950	4,7965 6,3739
	U/Th		0,7	1,0	1,7	1,0	5, L	2,2	1,6	4,8	4,8	1,3	2,3	4,4	0.7	1,8	2,2	1,7	1,4	38,2	3,3	1,6	8,1	2,7	11,1	2,9	1,2	2,4	3,7	1,5	1.8	4,1	0,8	20,3	1,8	0,4	3,0	4,9	1,1	0.8	6,5	9,3	27,7	1,0	1,3	1,9 19,7
	206Pb	204Pb	3326	12730	11324	7678	17682	71566	9646	44250	59038	22354	20958	36610	17462	16938	56202	25924	11978	93874	64826	17570	45274	19818	132662	27440	33466	54522	41134	15664	12246	51792	4166	88936	44726	4674	6610	34888	33060	15702	39954	76768	60146	12994	25992	14868 14568
	(mqq) U		125	409	248	139	174	485	64	325	516	499	166	274	169	141	817	218	92	1244	351	208	370	240	1988	193	321	442	386	81	63	482	31	747	233	27	221	311	204	290	349	513	1039	179	504	146
	Analysis		ELM06-PV10-27	ELM06-PV10-54	ELM06-PV10-73	ELM06-PV10-14 FI M06-PV10-15	ELM06-PV10-49	ELM06-PV10-71	ELM06-PV10-102	ELM06-PV10-31	ELM06-PV10-104	ELM06-PV10-62	ELM06-PV10-93	ELM06-PV10-13	ELM06-PV10-6	ELM06-PV10-7	ELM06-PV10-9	ELM06-PV10-97	ELM06-PV10-96	ELM06-PV10-48	ELM06-PV10-69	ELM06-PV10-80	ELM06-PV10-74	ELM06-PV10-8	ELM06-PV10-53	ELM06-PV10-11	ELM06-PV10-98	ELM06-PV10-67	ELM06-PV10-46	ELM06-PV10-56	ELM06-PV10-84	ELM06-PV10-39	ELM06-PV10-109	ELM06-PV10-55	ELM06-PV10-66	ELM06-PV10-51	ELM06-PV10-18	ELM06-PV10-92	ELMU0-PV10-88	ELM06-PV10-61	ELM06-PV10-30	ELM06-PV10-91	ELM06-PV10-33	ELM06-PV10-1	ELM06-PV10-17	ELM06-PV10-95 ELM06-PV10-25

	Age of youngest cluster (Ma)	(3+ analyses)																								155 (4)																			
	+1	(Ma)	17,8	23,1	1/'/	31.8	41,4	17,7	24,6	17,7	17,7	40,6	76,1	63,1 28.6	52,1	19,1	16,5	16,5	17,0	16,5	20,9	16,5	16.4	16,4	-		3,0	3,7	4,7	0,4 4,6	3,4	11,3	6,0	2,9	4,5	8,8	м, с 1, 8	2,2	3.0	6,5	8,4	16,2	6,8	3,3	12,4
-	Best age	(Ma)	1996,0	2003,2	2007,3	2009,8	2014,2	2017,2	2030,5	2033,8	2041,0	2062,5	2075,9	2095,7	2419,1	2662,0	2704,3	2721,1	2724,3	2733,4	2736,9	2755.4	763.0	2801,5			153,4	155,5	156,3	224.7	229,8	236,4	239,8	241,2	241,2	243,6	251,7	25/,4	258.9	267,0	269,3	270,2	274,8	281,1	282,7
-	+1	(Ma)	17,8	23,1	///	31.8	41,4	17,7	24,6	17,7	17,7	40,6	76,1	63,1 28.6	52,1	19,1	16,5	16,5	17,0	16,5	20,9	16,5	16.4	16,4			100,4	79,6	165,5	814,9 114,4	185,9	555,1	87,3	121,5	40,3	122,6	46,8	24,2 95.4	89.5	67,9	114,7	582,9	229,2	78,2	321,8 151,2
-	206Pb*	207Pb*	1996,0	2003,2	2007,3	2009,8	2014,2	2017,2	2030,5	2033,8	2041,0	2062,5	2075,9	2095,7	2419,1	2662,0	2704,3	2721,1	2724,3	2733,4	2736,9	2739,9	1/2/2	2801,5			61,5	82,4	91,4	110.6	58,6	52,9	293,8	260,5	241,1	122,0	215,8	304.8	171.2	202,4	178,7	46,8	60,4	204,9	137,1
jes (Ma)	+1	(Ma)	12,4	20,1	12,4	17,1	22,0	16,5	15,1	12,4	12,3	21,9	39,0	32,8	29,6	14,4	13,4	13,4	13,8	13,4	15,3	13,4	13.4	13,5			6,4	5,8	10,8	9.9	15,4	46,6	10,0	11,7	5,6	13,2	5,6	101	8,9	8,8	13,5	54,8	22,3	8,6	45,8
Apparent ag	207Pb*	235U	2037,4	2023,2	202/,2	2014,5	1940,4	2046,4	2067,2	2046,9	1975,0	1989,8	2028,8	2044,7	2351,2	2667,9	2698,8	2759,2	2729,4	2735,9	2739,5	27734 8	27817	2828,9			148,0	151,1	152,4	215,0	215,2	220,4	244,8	243,0	241,2	232,5	248,3	7 696	250.3	260,5	260,1	248,3	253,4	273,1	340,2 269,7
	+1	(Ma)	17,8	33,1	2'1/2	2//0	16,2	28,2	18,0	17,6	16,5	16,6	17,1	17,2	19,2	21,9	22,0	22,8	23,0	22,3	22,3	22,9	77.8	23,1			3,0	3,7	4,7	0,0 14	3,4	11,3	6,0	2,9	4,5	8,8	о 8 г 8	2,2 2,1	3.0	6,5	8,4	16,2	6,8	3,3	12,4
	206Pb*	238U	2078,6	2042,8	2046,/	1866.4	1871,9	2075,4	2104,1	2060,0	1912,5	1920,5	1982,9	1994,6 2147.0	2273,7	2675,5	2691,6	2811,5	2736,3	2739,2	2743,1	2825,3	2807.6	2867,6			153,4	155,5	156,3	224.7	229,8	236,4	239,8	241,2	241,2	243,6	251,7	758.0	258.9	267,0	269,3	270,2	274,8	281,1	282,7 285,3
-	error	corr.	0,71	0,82	1//0	0,82	0,39	0,85	0,58	0,71	0,71	0,40	0,23	0,27	0,31	0,66	0,71	0,71	0,71	0,71	0,62	0,71	0 71	0,71			0,42	0,58	0,40	0,11	0,19	0,21	0,56	0,22	0,74	0,58	0,61	0,09	0,30	0,64	0,55	0,25	0,25	0,34	0,28
-	+1	(%)	1,0	1,9	0, ,	1,0	1,0	1,6	1,0	1,0	1,0	1,0	1,0	1,0	1,0	1,0	1,0	1,0	1,0	1,0	1,0	0,1	- r	1,0			2,0	2,4	3,1 1	3,7	1,5	4,9	2,6	1,2	1,9	3,7	1,5	1,0	1.2	2,5	3,2	6,1	2,5	1,2	1,7
ratios	206Pb*	238U	0,3805	0,3728	0,3/3/	0,36/8	0,3369	0,3798	0,3860	0,3765	0,3454	0,3471	0,3602	0,3626	0,4229	0,5144	0,5182	0,5467	0,5288	0,5295	0,5304	0,5500	0 5458	0,5602			0,02408	0,02441	0,02454	0,03547	0,03629	0,03736	0,03789	0,03813	0,03813	0,03851	0,03982	0.040/4	0.04097	0.04229	0,04266	0,04280	0,04355	0,04456	0,04483
Isotopic	+1	(%)	1,4	2,3	1,4	2,0	2,5	1,9	1,7	1,4	1,4	2,5	4,4	3,7	3,2	1,5	1,4	1,4	1,5	1,4	1,6	1 4 1	- T	1,4,1			4,6	4,1	7,6	5.1 5.1	7,9	23,5	4,6	5,4	2,6	6,4	2,5	4,1	4.0	3,8	5,9	24,9	9,9	3,6	15,8 6,6
	207Pb*	235U	6,4376	6,3341	6,3629	6,2719 5,7386	5,7596	6,5036	6,6589	6,5077	5,9941	6,0966	6,3750	6,4914	9,1304	12,8383	13,2670	14,1409	13,7029	13,7975	13,8501	14,3899	144808	15,2168			0,15687	0,16040	0,16190	0,10289	0,23608	0,24247	0,27269	0,27043	0,26818	0,25736	0,27703	0,28013	0.27959	0,29248	0,29202	0,27706	0,28352	0,30854	0,39807 0,30429
	U/Th		2,0	1,7	4,2	11,8	1,2	2,2	2,8	3,7	2,9	4,0	1,0	1,2	10,2	2,0	2,6	1,2	1,0	1,6	2,5	2,0	2,4	3,6			0,8	4,1	0,6	-,4 0,4	0,8	0,6	0,7	0,5	2,0	0,3	0,7	0,3	0,9	1.7	0,5	0,5	0,3	1,1	0,5
	206Pb	204Pb	32768	34980	/1832	1286/6	2182	30720	37350	46748	27980	19930	8132	35400	42784	29358	25114	25714	16256	33446	12028	43974	56490	118846			8350	10930	2468	6128 6128	4229	1153	5726	4550	20193	2615	9190	6786	6338	6893	1763	526	2680	8702	1004 5564
	U (ppm)		181	215	409	801	237	190	181	239	340	110	54	181	283	186	158	115	74	126	62	175	205	884		lvoi Island	504	651	117	505	178	32	253	185	503	175	304	1710	239	292	74	23	122	312	102
-	Analysis		ELM06-PV10-83	ELM06-PV10-59	ELM06-PV10-110	ELM06-PV10-70 FI M06-PV10-81	ELM06-PV10-19	ELM06-PV10-106	ELM06-PV10-82	ELM06-PV10-68	ELM06-PV10-24	ELM06-PV10-57	ELM06-PV10-77	ELM06-PV10-20 FI M06-PV10-101	ELM06-PV10-22	ELM06-PV10-3	ELM06-PV10-41	ELM06-PV10-79	ELM06-PV10-40	ELM06-PV10-35	ELM06-PV10-16	ELM06-PV10-108 ELM06-PV10-21	FI M06-PV10-86	ELM06-PV10-45		Sample 53.1 Stobo	53102-21	53102-63	53102-56	53102-22	53102-48	53102-43	53102-105	53102-28	53102-68	53102-8	53102-80	53102-19 53102-17	53102-86	53102-33	53102-50	53102-85	53102-72	53102-100	53102-88 53102-108

	Age of youngest cluster (Ma)	(3+ analyses)																																												
	+1	(Ma)	5,8	5,9	5,6	12,3	5.1	3,6	4,6	41,7	25,8	29,9	31,2	27,2	18.1	26,5	26,2	34,3	20,6	41,8	18,2	18,1	34,1	39,0	26,9	18,0	28,6	18,2	22,1	24,4	33.1	19,2	26,2	29,9	25,3	23,6	30,1	26,7	26,3	17,9	37,8	45,0	33,6	27,7	48,3	21,5 41,3
-	Best age	(Ma)	287,5	290,6	291,8	292,0 295 0	309.5	319,2	330,8	1703,2	1762,4	1766,8	1838,7	1842,9	1857.6	1858,5	1860,1	1860,5	1863,1	1863,3	1869,4	1870,0	1872,6	1872,9	1886,2	1890,6	1891,5	1892,7	1894,5	1910,9	1916.2	1916,7	1923,4	1924,2	1929,6	1931,3	1932,2	1934,5	1941.5	1944,6	1950,8	1951,1	1959,6	1963,0	19/9/0	1991,7
-	+1	(Ma)	111,5	145,4	71,7	135,6	382.8	26,7	48,6	41,7	25,8	29,9	31,2	27,2	18,1	26,5	26,2	34,3	20,6	41,8	18,2	18,1	34,1	39,0	26,9	18,0	28,6	18,2	22,1	24,4	33.1	19,2	26,2	29,9	25,3	23,6	30,1	26,7	26.3	17,9	37,8	45,0	33,6	27,7	48,3	21,3 41,3
-	206Pb*	207Pb*	182,4	280,5	223,0	425,8	2.8	336,9	286,3	1703,2	1762,4	1766,8	1838,7	1842,9	1857.6	1858,5	1860,1	1860,5	1863,1	1863,3	1869,4	1870,0	1872,6	1872,9	1886,2	1890,6	1891,5	1892,7	1894,5	1910,9	1916.2	1916,7	1923,4	1924,2	1929,6	1931,3	1932,2	1934,5	1941.5	1944,6	1950,8	1951,1	1959,6	1963,0	19/9/0	1993,9
ges (Ma)	+1	(Ma)	12,6	16,8	9,1	19,7	38.5	4,6	7,1	20,6	15,4	32,0	19,5	68,4 24 F	12.4	15,2	19,2	19,3	19,7	22,1	12,1	13,7	28,7	22,8	19,8	26,8	21,0	17,3	19,6	20,1	34.5	12,7	16,0	17,7	17,7	19,0	18,5	22,5	25.6	14,9	39,3	28,1	22,5	27,9	33,8	22,8
Apparent a	207Pb*	235U	276,3	289,5	284,3	307,4	276.2	321,3	325,3	1724,4	1773,8	1594,4	1859,5	1656,5	1820.9	1879,5	1876,3	1851,8	1835,5	1866,8	1858,0	1855,3	1901,5	1881,7	1772,7	1923,2	1880,7	1900,8	1776,1	1932,8	1910.0	1932,6	1941,8	1926,0	1893,6	1761,1	1929,6	1941,3	1901.2	1967,7	1931,1	1972,3	2003,9	1978,7	1956,2	2012,1
	+1	(Ma)	5,8	5,9	5,6	12,3	5 <u>1</u>	3,6	4,6	15,5	18,3	47,5	24,4	36.0	16.8	16,4	28,1	19,9	31,8	18,7	16,1	20,1	45,8	25,5	27,1	49,5	30,4	28,8	29,3	31,7	58.7	16,8	19,3	19,9	24,3	26,4	22,1	35,8	42.0	23,7	66,6	34,5	30,8	48,1	46,/	45,6 20,7
-	206Pb*	238U	287,5	290,6	291,8	292,0 205 0	309.5	319,2	330,8	1741,8	1783,5	1467,2	1878,1	1513,5	1789.1	1898,6	1891,0	1844,0	1811,2	1869,8	1847,8	1842,3	1928,1	1889,7	1677,9	1953,7	1870,9	1908,1	1677,1	1953,3	1904.3	1947,4	1959,1	1927,7	1860,8	1621,2	1927,3	1947,7	1864.6	1989,8	1912,8	1992,5	2047,1	1993,8	1934,8	2029,7
	error	corr.	0,40	0,31	0,53	0,58	0.11	0,70	0,56	0,41	0,64	0,91	0,66	0,98	0.73	0,56	0,76	0,55	0,87	0,45	0,70	0,78	0,82	0,58	0,78	0,95	0,76	0,87	0,85	0,81	0,89	0,68	0,62	0,58	0,73	0,81	0,62	0,82	0.87	0,81	0,89	0,62	0,68	0,88	0,/2	0,91
	+1	(%)	2,1	2,1	2,0	4,3	2,1	1,2	1,4	1,0	1,2	3,6	1,5	8,2	1,1	1,0	1,7	1,2	2,0	1,2	1,0	1,3	2,7	1,6	1,8	2,9	1,9	1,7	2,0	1,9	3.6	1,0	1,1	1,2	1,5	1,8	1,3	2,1	2.6	1,4	4,0	2,0	1,8	2,8	2,8	1,2
ratios	206Pb*	238U	0,04561	0,04612	0,04630	0,04634	0.04918	0,05076	0,05266	0,31023	0,31872	0,25559	0,33821	0,26464	0.31986	0,34247	0,34090	0,33116	0,32441	0,33650	0,33195	0,33080	0,34865	0,34064	0,29731	0,35400	0,33673	0,34446	0,29714	0,35391	0.34367	0,35269	0,35513	0,34855	0,33464	0,28594	0,34846	0,35274	0.33542	0,36160	0,34544	0,36219	0,37377	0,36245	0,35004	0,37007
sotopic	+1	(%)	5,2	6,7	3,7	7,4	15.9	1,7	2,6	2,5	1,8	4,0	2,3	8,4 0 C	1.5	1,8	2,2	2,3	2,3	2,6	1,4	1,6	3,3	2,7	2,4	3,1	2,5	2,0	2,3	2,3	4.0	1,5	1,9	2,1	2,1	2,3	2,1	2,6	3.0	1,7	4,5	3,2	2,6	3,2	2,9	3,0 2,6
	207Pb*	235U	0,31276	0,32993	0,32308	0,35358	0.31262	0,37229	0,37768	4,46426	4,73686	3,80783	5,24192	4,11107	5,00957	5,36635	5,34653	5,19505	5,09617	5,28694	5,23295	5,21649	5,50592	5,38034	4,73074	5,64649	5,37392	5,50102	4,75000	5,70934 E 2E0EE	5.56068	5,70806	5,76906	5,66481	5,45508	4,66567	5,68862	5,76588 5,16621	5.50412	5,94417	5,69826	5,97565	6,19617	6,01985	5,86608	5,94257 6,25415
	U/Th		0,4	0,9	1,1	0,6	<u>, -</u>	1,6	1,0	1,0	1,1	1,0	0,9	2,3	01	0,8	0,7	1,9	1,6	1,1	1,2	1,7	1,2	1,4	1,8	1,2	1,7	1,3	2,3	7,0	0,6	1,5	3,6	0,4	1,8	2,8	2,1	3,2	3.0	2,7	2,5	1,8	2,6	4,6	0,1	1,5 10,3
	206Pb	204Pb	4194	2478	4448	1422	2647	11990	9235	4813	20617	6559	12896	81237	35656	16175	13702	89390	36870	9611	17690	23846	29910	10883	9350	39190	15064	90369	156188	256159	19820	71930	105207	52652	59204	145229	40618	54154	5810	65004	17029	95936	34977	107778	8//45	54971 101295
	(mqq) U		152	94	200	1031	02	650	269	22	102	249	62	373	153	65	65	348	148	33	61	187	298	35	41	153	59	447	869	1706	123	236	457	204	249	860	176	152	82	328	476	502	112	481	462	231 651
	Analysis		53102-31	53102-97	53102-83	53102-24 53102-60	53102-92	53102-53	53102-11	53102-14	53102-49	53102-55	53102-26	53102-71	53102-91	53102-44	53102-81	53102-90	53102-3	53102-30	53102-93	53102-27	53102-95	53102-66	53102-102	53102-61	53102-45	53102-34	53102-62	53102-78	53102-2	53102-35	53102-89	53102-42	53102-96	53102-57	53102-64	53102-67	53102-98	53102-51	53102-10	53102-15	53102-75	53102-54	53102-104	53102-58 53102-60

			-	sotopic r	atios					Apparent a	ges (Ma)					
(mq	206Pb	U/Th	207Pb*	+	206Pb*	+1	error	206Pb*	+	207Pb*	+	206Pb*	+I	Best age	+1	Age of youngest cluster (Ma)
	204Pb		235U	(%)	238U	(%)	corr.	238U	(Ma)	235U	(Ma)	207Pb*	(Ma)	(Ma)	(Ma)	(3+ analyses)
67	18102	1,3	5,89279	1,6	0,34587	1,2	0,77	1914,8	20,1	1960,2	13,7	2008,4	17,8	2008,4	17,8	
76	19539	3,2	6,60056	2,2	0,38565	1,6	0,73	2102,6	28,8	2059,4	19,4	2016,4	26,8	2016,4	26,8	
240	27937	0,4	6,46070	1,5	0,37273	1,0	0,69	2042,2	18,1	2040,6	13,2	2038,9	19,3	2038,9	19,3	
672	93540	1,6	6,38039	2,0	0,36608	1,2	0,63	2010,9	21,4	2029,6	17,2	2048,6	26,9	2048,6	26,9	
43	12993	7,6	6,68335	6,2	0,38017	5,5	0,89	2077,1	97,6	2070,4	54,8	2063,8	50,7	2063,8	50,7	
678	132740	1,9	6,86487	1,9	0,37895	1,5	0,78	2071,4	26,0	2094,1	16,8	2116,6	20,9	2116,6	20,9	
676	17942	1,0	7,07802	4,0	0,37287	3,7	0,91	2042,9	64,0	2121,3	35,7	2198,1	28,6	2198,1	28,6	
368	103687	1,6	8,90521	2,2	0,42827	1,9	0,83	2297,9	35,8	2328,3	20,2	2355,1	20,8	2355,1	20,8	
233	74128	1,4	8,37517	2,5	0,39458	2,3	0,92	2144,0	41,5	2272,5	22,5	2390,2	17,0	2390,2	17,0	
91	29357	1,8	10,11424	2,6	0,44538	2,4	0,92	2374,7	46,8	2445,3	23,7	2504,5	16,8	2504,5	16,8	
116	19240	1,3	9,95350	2,7	0,42896	2,1	0,75	2301,0	40,0	2430,5	25,4	2540,7	30,4	2540,7	30,4	
796	54734	43,0	12,70687	2,6	0,51708	2,1	0,79	2686,8	45,4	2658,2	24,5	2636,5	26,2	2636,5	26,2	
309	5550	4,1	10,93319	4,8	0,44235	4,1	0,85	2361,2	80,4	2517,5	44,7	2646,0	42,3	2646,0	42,3	
589	81286	1,6	12,18589	4,4	0,49198	4,2	0,96	2579,2	89,7	2618,8	41,2	2649,5	20,1	2649,5	20,1	
1074	34581	2,5	12,12039	2,2	0,48896	1,9	0,87	2566,2	40,6	2613,8	20,8	2650,8	18,5	2650,8	18,5	
302	64450	1,5	12,40643	5,9	0,49829	5,5	0,93	2606,4	118,5	2635,7	55,9	2658,2	36,3	2658,2	36,3	
31	11907	0,6	12,86853	2,7	0,51238	2,5	0,93	2666,8	55,6	2670,1	25,8	2672,5	16,6	2672,5	16,6	
46	16422	0,7	13,30634	2,1	0,52644	1,2	0,54	2726,4	25,8	2701,6	20,2	2683,1	29,8	2683,1	29,8	
98	24742	0,4	12,47779	2,5	0,49315	1,9	0,77	2584,3	40,5	2641,1	23,2	2684,8	26,0	2684,8	26,0	
103	32250	0,9	12,47514	4,7	0,49200	3,7	0,77	2579,3	77,7	2640,9	44,5	2688,4	49,8	2688,4	49,8	
96	38135	0,6	13,87382	4,5	0,53863	4,1	0,91	2777,7	93,1	2741,1	43,0	2714,3	31,3	2714,3	31,3	
259	78208	1,5	13,72069	1,7	0,53172	1,0	0,59	2748,7	22,4	2730,6	16,1	2717,3	22,7	2717,3	22,7	
360	94599	1,2	13,05797	2,0	0,50295	1,5	0,76	2626,4	33,1	2683,8	19,2	2727,4	21,9	2727,4	21,9	
199	46612	1,5	13,65921	2,1	0,52233	1,8	0,83	2709,1	39,3	2726,4	20,3	2739,2	19,7	2739,2	19,7	
136	59081	1,3	14,52497	3,7	0,55325	3,1	0,85	2838,7	71,7	2784,6	34,8	2745,7	31,6	2745,7	31,6	
166	26971	1,1	14,03611	2,4	0,53301	2,1	0,86	2754,1	47,0	2752,1	23,1	2750,7	20,4	2750,7	20,4	
706	43750	1,0	12,97117	9,3	0,47747	5,6	0,61	2516,2	117,0	2677,6	87,6	2801,7	120,7	2801,7	120,7	
231	72870	1,5	15,31358	1,9	0,54370	1,4	0,72	2798,9	30,8	2834,9	17,9	2860,6	21,2	2860,6	21,2	
190	36728	4,1	16,26024	3,3	0,54223	2,9	0,89	2792,8	66,3	2892,2	31,5	2962,2	24,2	2962,2	24,2	
2418	154385	2,2	17,80254	9,3	0,53014	7,3	0,79	2742,0	163,2	2979,1	89,3	3143,3	90,6	3143,3	90,6	
673	9520	1,3	19,84068	2,4	0,58529	2,2	0,91	2970,3	53,1	3083,6	23,6	3158,3	15,9	3158,3	15,9	

All errors are shown at the 1-sigma level, and include only measurement errors. Systematic errors (mainly from fractionation correction) add ~1% uncertainty (2-sigma) to all ages. U concentration and U/Th have uncertainties of ~25%. 206Pb/204Pb is measured ratio. Decay constants: 235U=9.8485x10-10. 238U=1.55125x10-10, 238U/235U=137.88. Isotope ratios are corrected for Pb/U fractionation by comparison with standard zircon with an age of 564 \pm 4 Ma (2-sigma). Initial Pb composition interpreted from Stacey and Kramers (1975), with uncertainties of 1.0 for 206Pb/204Pb and 0.3 for 207Pb/204Pb.



Figure 5. Cumulative age probability curves (e.g. Gehrels 2006) and relative probability distribution diagrams (inset) for detrital zircon ages.

samples (Fig. 5) and the relative age probability distribution diagrams for detrital grains younger than 1 Ga allows a more detailed comparison of the younger populations of zircons in the samples (Fig. 6).

All of the samples from Stolbovoi and Chukotka have a few Archean grains, but the age distributions have a high percentage of 2.1-1.7 Ga zircons (44-67% of the populations). This fact, together with the arkosic and mica-rich nature of the sandstones can be used to infer that Precambrian crystalline rocks were likely proximal sources for the sediments. Essentially no zircons between the ages of 1700 and 500 Ma are present in any of the samples. A second set of ages, spanning the Late Paleozoic (~320-250 Ma) is the next most abundant zircon age population. Despite the abundance of what



Figure 6. Comparison of detrital zircon populations younger than 1 Ga from syn-orogenic Jurassic-Cretaceous sandstones of the Russian Arctic. Light grey line represents combined data from three Triassic samples previously described from Chukotka (Miller et al. 2006; all LA-ICPMS ages, all grains < 1000 Ma, N=241).

appear to be lithic fragments of older Triassic shales and sandstones in the Jurassic-Cretaceous deposits and the fact that Triassic strata have been found to contain zircon populations of about this age (Fig. 6, shown by grey line, from Miller et al. 2006) it is unlikely that the Late Paleozoic zircons in these rocks are mostly recycled. The reason for this is that the Triassic sediments previously studied also contain abundant Early Paleozoic zircons, which are not present in significant abundance in the younger Jurassic-Cretaceous strata (Fig. 6). It would be very difficult to recycle Late Paleozoic zircons from Triassic rocks into Jurassic-Cretaceous strata without also recycling some of their Early Paleozoic zircons. Based on this reasoning, we infer that Late Paleozoic magmatic rocks constituted part of the source region for Jurassic-Cretaceous sandstones. A smaller percent of detrital zircons in the Jurassic-Cretaceous sediments (~3-9%) are Jurassic in age, spanning ~175-145 Ma, with their youngest ages approaching the (fossil-dated) depositional age of the sandstones (Figs. 5 and 6, Table 3). Given the large percent of volcanic lithic fragments in the sandstones, together with the ages of detrital zircons, it seems reasonable to infer that both Late Paleozoic and Mesozoic volcanic rocks may have also been common rocks in the source regions, but likely contributed fewer zircons to the populations as compared to Precambrian gneisses and granitic rocks and Late Paleozoic and Mesozoic plutons.

The distributions of U-Pb detrital zircon ages in all of these sandstones are very similar to nearly identical (Figs. 5 and 6) despite the present distances between samples (Table 2). This is even more remarkable given the fact that the exact age of deposition (thus relative stratigraphic position) of the samples is not exactly known (due to lack of fossils, poor exposure, complex structure and large distances between the outcrops sampled). The high P values from the K-S test (Table

	est results for JU	Jrassic-Cretace	eous samples r	rom the Russic	In Arctic.	
		K-S P-values using	g error in the CDF			
	53/1-02	GB9986	CH04ELM7	04JT54C	ELMCH03 24.3	ELM06-PV10
53/1-02		0.813	0.221	0.200	0.103	0.459
GB9986	0.813		0.668	0.809	0.081	0.094
CH04ELM7	0.221	0.668		0.999	0.025	0.021
04JT54C	0.200	0.809	0.999		0.010	0.006
ELMCH03 24.3	0.103	0.081	0.025	0.010		0.097
ELM06-PV10	0.459	0.094	0.021	0.006	0.097	

Table 4. K-S	S test results fo	r Jurassic-Cretaceous	samples from	the Russian	Arctic.
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Notes: The K-S test is a non-parametric method for comparing cumulative probability distributions. P(KS) gives the probability that random chance alone might produce the observed difference in two distributions drawn from the same parent population. A low probability on the test, such as P(KS) < 0.05, would indicate that the differences between the two distributions are significant and that the samples are not similar in terms of their age population If P(KS) >> 0.05 then the differences are just a factor of random chance. To apply the K-S test to the data, we used the algorithm of Guynn (2006). Values that pass the K-S test at 95% confidence level and are not rejected are shown in bold letters and highlighted in yellow.

4) confirm the noted similarities between samples. ELMCH03 24.3 and ELM06 PV 10 differ the most only because of somewhat different proportions of particular age populations (Figs. 5 and 6). They have less Jurassic zircons, for instance, but are the northernmost two samples from Chukotka, and lay further from the source region during deposition. The sample from the South Anyui Zone and the one from Stolbovoi Island are remarkably similar given that they are the furthest apart today, but both represent sites that were proximal to the SAZ. Their similarities are even more remarkable considering that their petrography indicates proximal derivation, ruling out extensive reworking and/or great distances of longitudinal transport within the basin.

Discussion of source regions for sandstones

In general, the involvement of a volcanic arc is implied in collisional models for the formation of the SAZ and the Chukotka-Anyui fold belt (e.g., Parfenov 1991; Sokolov 2002 and references therein), but the age range of magmatic activity in this arc is poorly known. Late Paleozoic intermediate to silicic volcanic sequences dated by fossils as Mississippian and Permian have been recognized in the South Anyui suture zone (Sokolov et al. 2006 and references therein). Jurassic to Early Cretaceous volcanic rocks also constitute an important structural component of the SAZ proper (the Jurassic Anyui-Svyatoi Nos island arc S of the New Siberian Islands (Kuzmichev et al. 2006)) and the Jurassic to Cretaceous Alazaya-Oloi arc of Chukotka (Dovgal 1964; Sokolov et al. 2002; Bondarenko et al. 2003). An extensive belt of granitoid plutons exists in northern Verkoyansk, south of the New Siberian Islands (Fig. 1B), called the "Main Belt" and has recently been dated by the U-Pb method utilizing the SHRIMP-RG. The intrusive history of this belt spanned the ~ 158-147 Ma time interval (Prokopiev et

al. 2007; Toro et al. 2007; Akinin et al. in press). This belt of plutons and/or silicic eruptions of this same age most likely provided the youngest zircons to the foreland basin deposits of the New Siberian Islands and the Myrgovaam-Raucha basin of Chukotka. Its location is not far from the New Siberian Islands but it lies considerably west of the Jurassic-Cretaceous deposits of Chukotka (Fig. 1A) and no major plutonic belts of this age are mapped south of the SAZ in Chukokta. There are granites mapped between the two regions that constitute what is called the "Northern Belt" but these are somewhat younger (~ 135-127 Ma, Prokopiev et al. 2007; Toro et al. 2007; Akinin et al. in press) and slightly younger than the bulk of the sedimentary succession we have studied (Tithonian to Valanginian, ~ 151-136 Ma (Gradstein et al. 2004)).

The detrital zircon age data presented here specify the age of Paleozoic and Mesozoic igneous rocks and confirms their involvement in the orogenic event that produced the Chukotka fold belt and which led to the formation of the SAZ. Both Late Paleozoic (~330-250 Ma) and Mesozoic (~175 to 145 Ma) igneous rocks are involved in the collisional event. The new data also provide better constraints on the time of closure of the South Anyui Ocean. This closure must have occurred prior to the Tithonian, when zircons were transported from southerly sources onto the AACM plate, across the SAZ. Our data further document that the orogenic highlands that shed debris into the Jurassic-Cretaceous syn-orogenic basin contained Precambrian crystalline rocks about 2.1- 1.7 Ga, an age characterizing the Siberian craton. This source region is somewhat more mysterious as there are no widespread exposures of crystalline rocks mapped directly within or south of the SAZ. However, Precambrian crystalline rocks underlie the broad Kolyma-Omolon region to the south of the SAZ (see discussion in Kolesov & Stone 2002) and are inferred to form the basement of at least part of the Main Belt granitoids, as some of these Mesozoic plutons contain inherited Precambrian cores. (Prokopiev et al. 2007; Toro et al. 2007; Akinin et al. in press). The thrust plates that carried Precambrian crystalline rocks to the surface in order to be eroded into the Myrgovaam foreland Basin may have been subsequently buried by younger sediments.

Discussion and conclusions

The data presented here helps establish that the New Siberian Islands and Chukotka, two distant regions, contain exposures of the same syn-orogenic foreland basin sequence, providing robust evidence for inclusion of both of these parts of Arctic Russia in the AACM as previously proposed (e.g. Parfenov & Natali'in 1985; Kuzmichev et al. 2006) (Fig. 1). However, the data present problems for the rotational opening model for the Arctic, because they further establish with greater certainty the large dimensions of this plate (Fig. 1). And certainly the Jurassic-Cretaceous match described here makes it impossible to break the plate in the center and rotate only the Chukotka part as suggested by Rowley & Lottes (1988). The discovery of Kuzmichev & Pease

(2007) that mafic dikes and sills in the New Siberian Islands link the New Siberian Islands to Siberia makes it even more difficult to rotate any part of the AACM with respect to Siberia. Our previous results tentatively linked Chukotka to NE Russia in Triassic times (Miller et al. 2006) and the new set of data presented here provides an even stronger argument for linking Chukotka to NE Russia, making it increasingly difficult to restore the AACM plate back to North America (Fig. 1B). The southerly sources documented in the syn-orogenic Jurassic-Cretaceous foreland basin sandstones could not have delivered sediment across a vast Anyui Ocean to the AACM if this ocean, or even part of it, still existed in Tithonian to Valanginian times. These relations thus tie the New Siberian Islands and Chukotka to Siberia by at least Tithonian times. Thus the Anyui Ocean must have finished closing by the latest Jurassic. These arguments make the rotational opening model for the Amerasian Basin, inferred to have taken place at a younger time 135-120 Ma (e.g., Lawver et al. 2002), increasingly difficult to champion.

To better explain the relations discussed above, we suggest an alternative hypothesis for the initial opening of the Amerasian Basin. This is depicted in simplified



Figure 7. Speculative diagram showing a possible explanation for the separation of syn-orogenic Jurassic-Cretaceous sediments (and their identical sources) to their present positions by the rift opening of the Makarov Basin and ~100% extension of the Siberian Shelf in an E-W direction north of the SAZ. The SAZ may have operated as a transform boundary separating a more extended region to the north from a less extended one to the south. It would have also served as the locus of deformation that displaced Paleozoic and Mesozoic arc complexes southwards towards the Pacific Ocean margin.

fashion in Figure 7. The hypothesis revives previous ideas that the Makarov Basin opened by rifting in an orthogonal direction from the Lomonosov Ridge (e.g. Sweeney et al. 1982; Taylor et al. 1981; Vogt et al. 1982). The interpretation is that of the first author and is not necessarily shared by all of us (see, for instance Kuzmichev 2009, for alternative models). It by no means addresses the complete history of the formation of the Amerasian Basin, which still remains a formidable question.

The model explains the remarkable similarity in Jurassic-Cretaceous sandstones in the New Siberian Islands to those of Chukotka (as well as good matches between older geologic units (e.g. Miller et al. 2006) by the rift separation of Chukotka from the New Siberian Islands during the opening of the Makarov Basin on the Amerasian side of the Lomonosov Ridge (Fig. 7). The present separation between these two regions with identical inferred sources is explained as a function of stretching of continental lithosphere beneath both the East Siberian Shelf and the low-lying coastal region of Arctic Russia, similar to that taking place in the Laptev Sea today, south of the Gakkel mid-ocean ridge (Fig. 1A) (Drachev et al. 1998; Franke et al. 2004). The observed separation could be accommodated by ~100% extension distributed across this broad region, decreasing original distances between these exposures from ~1400 to ~700 km or less, a degree of stretching that can be anticipated in rifted and stretched continental crust and along rifted continental margins. Although there is very little data, present-day crustal thicknesses beneath the East Siberian Shelf sedimentary sequences based on a single seismic line are about 15-25 km, with an overlying sedimentary sequence up to about 8 km thick on the shelf edge (Franke et al. 2004). Since we don't know if and how much this crust was magmatically added to during rifting, original crustal thicknesses may have been only 30- 50 km or so. We do know that the time of the hypothetical rifting that formed the basement to the East Siberian Shelf must post-date deposition (and deformation) of the syn-orogenic Jurassic-Cretaceous sediments described here (post 136 Ma, Gradstein et al. 2004). Widespread magmatism represented by plutons that range in age from 121 to 106 Ma in the New Siberian Islands and 116.9± 2.5 to 108.1±1.1 Ma in Chukotka (Figs. 2 and 3) may be coeval with rifting as they are known to have been emplaced during E-W to NW-SE extension as documented by hundreds of dike orientations (Miller & Verzhbitsky in press).

This hypothesis (Fig. 7) would imply that the SAZ is not what it has typically been viewed as: the remains of an extensive ocean basin that closed during rotational opening of the Amerasian Basin (the Anyui Ocean in the rotational model of Lawver et al. 2002, depicted in Fig. 1B). Rather, the SAZ in this model represents parts of an allochthonous belt of oceanic and arc rocks previously thrust onto the margin of NE Russia, subsequently slivered in a right-lateral direction along a broad, transform fault system that helped to accommodate a greater amount of extension to the north than to the south, and translated Chukotka far to the east and south (Fig. 7).

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Appendix I. U-Pb geochronologic methods

U-Pb geochronology of zircons was conducted by laser ablation multicollector inductively coupled plasma mass spectrometry (LA-MC-ICPMS) at the Arizona LaserChron Center (Gehrels et al. 2006, 2008). The analyses involve ablation of zircon with a New Wave DUV193 Excimer laser (operating at a wavelength of 193 nm) using a spot diameter of 35 microns. The ablated material is carried in helium into the plasma source of a GVI Isoprobe, which is equipped with a flight tube of sufficient width that U, Th, and Pb isotopes are measured simultaneously. All measurements are made in static mode, using Faraday detectors with 10e11 ohm resistors for ²³⁸U, ²³²Th, ²⁰⁸Pb, and ²⁰⁶Pb, a Faraday detector with a 10e12 ohm resistor for ²⁰⁷Pb, and an ion-counting channel for ²⁰⁴Pb. Ion yields are ~1.0 mv per ppm. Each analysis consists of one 12-second integration on peaks with the laser off (for backgrounds), 12 one-second integrations with the laser firing, and a 30 second delay to purge the previous sample and prepare for the next analysis. The ablation pit is ~12 microns in depth.

For each analysis, the errors in determining $^{206}Pb/^{238}U$ and $^{206}Pb/^{204}Pb$ result in a measurement error of ~1-2% (at 2-sigma level) in the $^{206}Pb/^{238}U$ age. The errors in measurement of $^{206}Pb/^{207}Pb$ and $^{206}Pb/^{204}Pb$ also result in ~1-2% (at 2-sigma level) uncertainty in age for grains that are >1.0 Ga, but are substantially larger for younger grains due to low intensity of the ^{207}Pb signal. For most analyses, the cross-over in precision of $^{206}Pb/^{238}U$ and $^{206}Pb/^{207}Pb$ ages occurs at ~1.0 Ga.

Common Pb correction is accomplished by using the measured ²⁰⁴Pb and assuming an initial Pb composition from Stacey & Kramers (1975) (with uncertainties of 1.0 for ²⁰⁶Pb/²⁰⁴Pb and 0.3 for ²⁰⁷Pb/²⁰⁴Pb). Our measurement of ²⁰⁴Pb is unaffected by the presence of ²⁰⁴Hg because backgrounds are measured on peaks (thereby subtracting any background ²⁰⁴Hg and ²⁰⁴Pb), and because very little Hg is present in the argon gas (background ²⁰⁴Hg = ~300 CPS).

Inter-element fractionation of Pb/U is generally \sim 20%, whereas fractionation of Pb isotopes is generally <2%.

In-run analysis of fragments of a large zircon crystal (generally every fifth measurement) with known age of 564 \pm 4 Ma (2-sigma error) is used to correct for this fractionation. The uncertainty resulting from the calibration correction is generally 1-2% (2-sigma) for both ²⁰⁶Pb/²⁰⁷Pb and ²⁰⁶Pb/²³⁸U ages.

Concentrations of U and Th are calibrated relative to SRM 610 trace element glass, which contains ~460 ppm of each element. The homogeneity of this glass makes it a better concentration standard than a natural zircon crystal, and the U/Th fractionation is similar for glass and zircon.

The analytical data are reported in Table 3. Uncertainties shown in these tables are at the 1-sigma level, and include only measurement errors.

Interpreted ages are based on $^{206}\text{Pb}/^{238}\text{U}$ for <1000 Ma grains and on $^{206}\text{Pb}/^{207}\text{Pb}$ for >1000 Ma grains. This division at 1000 Ma results from the increasing uncertainty of $^{206}\text{Pb}/^{238}\text{U}$ ages and the decreasing uncertainty of $^{206}\text{Pb}/^{208}\text{Pb}$ ages as a function of age. Analyses that are >20% discordant (by comparison of $^{206}\text{Pb}/^{238}\text{U}$ and $^{206}\text{Pb}/^{207}\text{Pb}$ ages) or >5% reverse discordant are not included.

Cumulative and normalized relative age-probability diagrams were generated and K-S analyses were performed using routines available from www.geo. arizona.edu/alc.

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